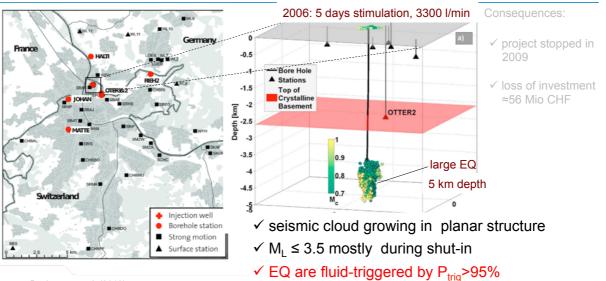
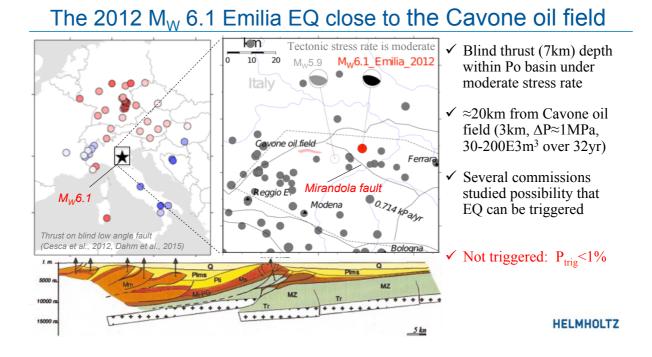


3rd Induced Seismicity Workshop, 5-9 March 2019 Schatzalp, Davos, Switzerland

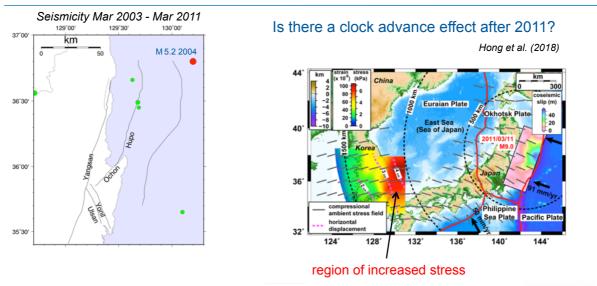


## The 2006 $M_L$ 3.4 Basel induced by deep hot rock stimulation

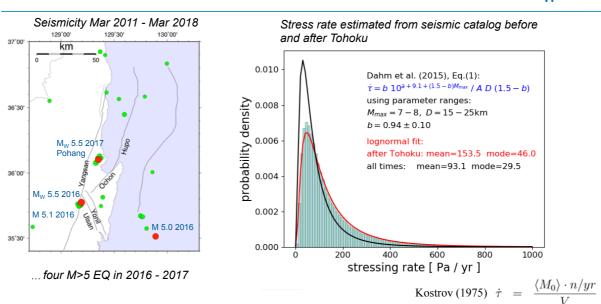
e.g. Bachmann et al. (2012)



S-Korea – far-distance stress shadow from 2011 Tohoku M<sub>w</sub> 9 ?

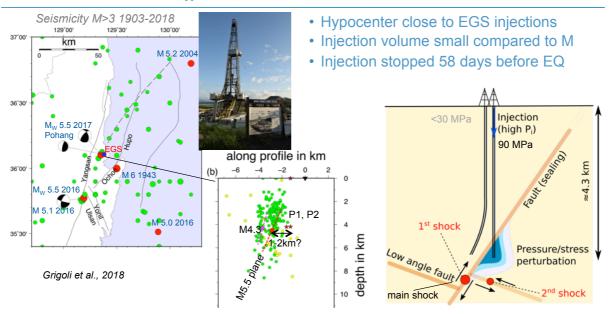


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# S-Korea – far-distance stress shadow from 2011 Tohoku $M_W$ 9?

#### Pohang 2017 M<sub>w</sub> 5.5 – induced, triggered or natural ?

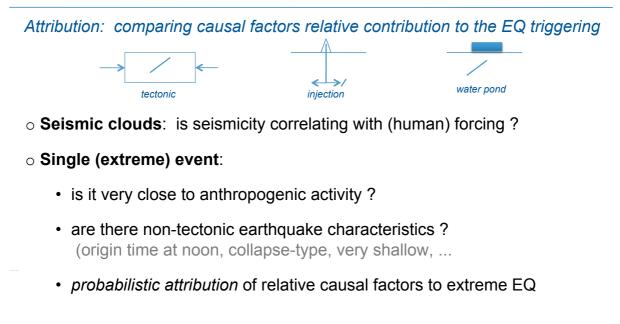


Pohang 2017 – a key test for attribution models

- ① Tectonic stress rate increased after the Tohoku 2011
- ② Pohang EQ is shallow and close to injection operation (but "too large" and 58 days after shut-in)

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# **Discrimination:** approaches



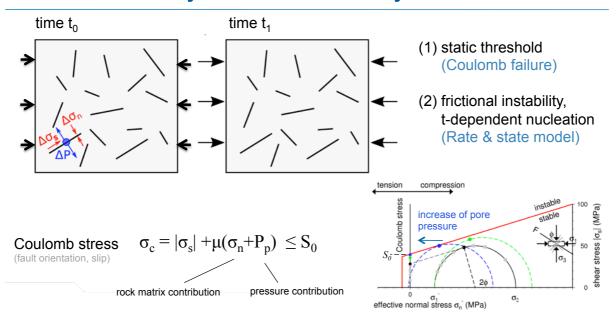
#### **Extreme event probabilistic attribution**

**Causal factors**:(1) tectonic Coulomb stress rate do<sub>c</sub>/dt

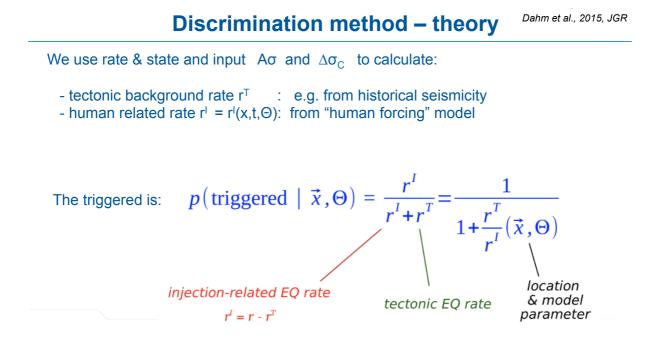
(2) anthropogenic stressing

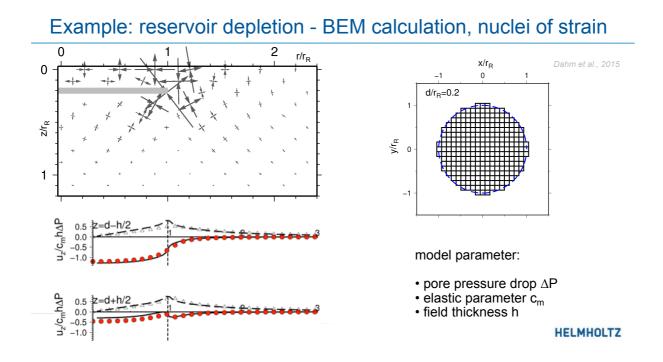
Probabilistic attribution (adapted from climate research):

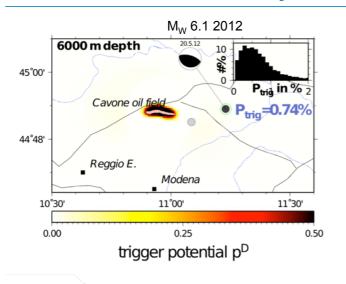
- a) physics-based seismicity model to assess relative contribution
  - (1) theoretical EQ rate  $r^{T}$  from tectonic stressing
  - (2) theoretical EQ rate r<sup>I</sup> from human action
- b) assigning **statistical confidence** that the EQ was human-triggered



#### **Physics-based seismicity model**



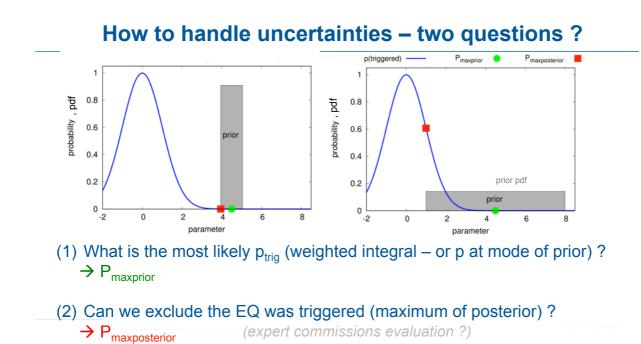


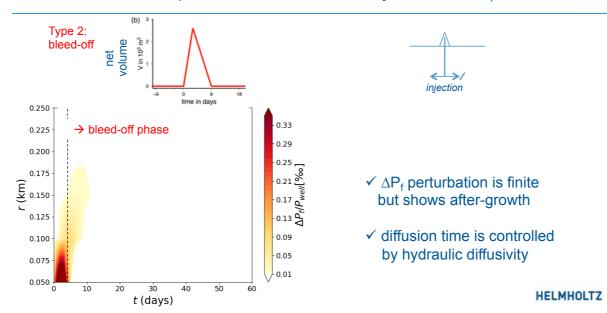


### Case study: Emilia oil field

P<sub>trig</sub> = < 1%

Dahm et al., 2015

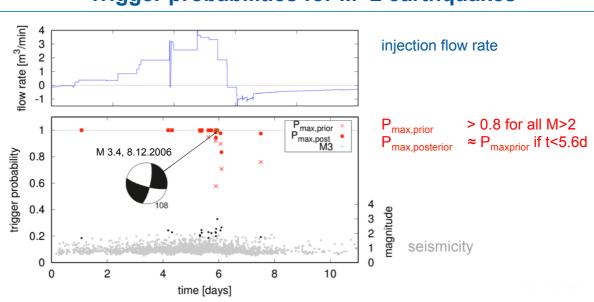




### Example: short-term water injection – $\Delta P_f$ diffusion

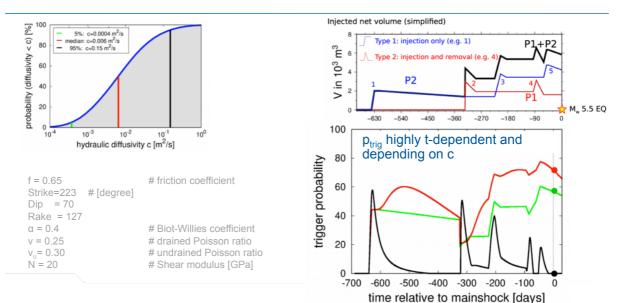
# Case study: 5 days stimulation beneath Basel

	0040 -	Dahm et al. (2015), Eq.(1):	(1) tectonic	(3) assumed parameter uncertainties			
density	030 -	$\dot{\tau} = b \ 10^{a+9.1+(1.5-b)M_{max}} / A D \ (1.5-b)$	stressing rate	Parameter	Mode	$\left[min,max\right],\pm std$	D
0.0 de	0025 -			EQ relative location SED catalogue			
probability o	0015 -			x, z, z in $[km]$	SED catalog	$\pm 0.1$	n
	0010 -	lognormal fit: mean=243.7 mode=115.3	EQ mechanism				
				see Deichmann & Giar	rdini (2009)	$\pm 10^{\circ}$	n
	stressing rate [Pa/yr] (2) diffusivity from			Rock and fluid parameter			
	800	D = 0.01 m <sup>2</sup> /s	growth of seismic cloud	Frict. coeff. f	0.68	[0.5, 0.85]	u
	700	D = 0.03 m <sup>2</sup> /s		Biot-Willies $\alpha$	0.5	[0.1, 0.9]	u
	600	D = 0.05 m <sup>2</sup> /s		Skempton $\alpha$	0.5	[0.1, 0.9]	u
	600			Poisson $\nu$ (drained)	0.245	[0.20, 0.29]	u
-	500	-		Poisson $\nu_u$ (undr)	0.3	[0.29, 0.31]	u
2		000		Rigidity $\mathcal{N}$ [GPa]	17.5	[5,30]	u
distance [m]	400			Fluid viscosity $\eta \left[ Pas \right]$	$3.5 \cdot 10^{-4}$	$[3,4] \cdot 10^{-4}$	u
dist	300			Permeability $\kappa [mD]$	0.1	[0.05, 0.5] 90% conf.	lg
	200			Diffusivity c $\left[m^2/s\right]$	0.03	s±0.01	
	100			Seismicity model parameter			
	0			frict. value $A\sigma$ [MPa]	0.03	[0.01, 0.05]	u
	(	0 2 4 6 time [days]	8 10	tect. stress $\dot{\tau} \left[ Pa/yr \right]$	186.7	see figure	lg



#### **Trigger probabilities for M>2 earthquakes**

Example: the late M<sub>w</sub> 5.5 Pohang earthquake



Summary of case studies											
Event	Pmaxprior	P <sub>maxposterior</sub>	Reference								
- quasi-static reservoir depletion -											
M <sub>w</sub> 4.3 2001 Ekofisk M <sub>w</sub> 4.4 2004 Rotenburg M <sub>w</sub> 6.1 2012 Emilia	>99% ≈70% <1%		Dahm et al. (2015) Dahm et al. (2015) Dahm et al. (2015)								
- time-dependent "injection" -											
$M_W$ 6.1 1976 Tjörnes fracture zon $M_W$ 3.4 2006 Basel	ne >90% 99%	- >95%	Passarelli et al. (2011)								
M <sub>w</sub> 5.5 2017 Pohang			in prep.								

Probabilistic attribution approach is flexible to be applied to different problems

### Conclusion

 ✓ Expert panel reports usually assess the P<sub>maxposterior</sub> (assess whether triggering can be excluded assuming it was triggered)

✓ The likelihood to be triggered (mode, P<sub>maxprior</sub>) is smaller than P<sub>maxposterior</sub>

 $\checkmark$  We suggest to report both measures, as purely statistical bounds are more difficult to communicate