Fluid-Induced Seismicity – Comparison of Rate- and Stateand Critical Pressure Theory

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* Abstract Induced seismicity as generated by the injection of fluid in a homogeneous, porous medium with faults with variable proximity to rupture conditions is simulated using rate- and state-dependent frictional fault theory (RST) of Dieterich (1994) and the critical pressure theory (CPT) developed by Shapiro (2015). In CPT the seismicity is proportional to the pressure rate but limited by the Kaiser Effect. There is no time delay between a change in pressure rate and change in seismicity. RST is more complex and includes a time delay between a pressure change and the seismicity. Comparing both modelling approaches at fixed location this delay can be significant. However, it is small where the seismicity is high and larger where it becomes small. The evolution of the total seismicity in the medium where fluid is injected with time is thus very similar with both modelling approaches so that RST and CPT provide very similar results.

Section 2018 Section 2018 Essentials of the Critical Pressure **Theory (CPT)**

fault density: $n_f(\vec{x})$ probability that the criticality parameter is less than C: $F_c(C)$ pressure due to injection: $p(\vec{x}, t)$ Seismicity rate:



$$N(t) = \int n_f(\vec{x}) \cdot f_c(p(\vec{x}, t)) \cdot \frac{\partial p(\vec{x}, t)}{\partial t} \cdot dV$$

 $\frac{\partial p(\vec{x},t)}{\partial t}$ must be positive and the Kaiser Effect considered. For constant volume rate and homogeneous diffusion

$$p(r,t) = \frac{\dot{Q_0}}{4\pi DS_r} \cdot \left\{ 1 - erf\left(\frac{r}{\sqrt{4Dt}}\right) \right\}$$

Seismicity rate is:

a 8

Injection Pressure with Time

$$N(t) = \frac{n_f}{C_2 - C_1} \cdot \frac{\dot{Q_0}}{S} = N_0$$

Fig. 2: Pressure and temporal pressure derivative at 30m distance, scaled to relative amplitudes so that they fit into one plot. In the left panel the derivative is set to zero only if it becomes negative. The right panel considers the Kaiser Effect and keeps $\dot{p}(r,t)$ at zero level until the pressure in the second cycle exceeds the maximal value it experienced before.



time in



 $b(t) = \frac{\dot{\tau}_{tec}}{C_1} \cdot \left\{ 1 + \frac{p(t)}{\sigma_n^0} \right\}$ $a(t) = b(t) + \frac{g_0}{C_1} \cdot \dot{p}(t)$

Fig. 4: Solution of Ricatti Equation (11) at distance 30 m (left) and comparison with the CPT ansatz that includes the Kaiser Effect (right).



Fig. 1: Left panel shows injection pressure (over hydrostatic) at the origin with two cycles of about 14 h each and 10 MPa amplitude. The right panel show pressure at 30, 50, and 100m calculated with a diffusion constant of $D = 0.1m^2/s$.

The Rate- and State-dependent **Theory (RST)**

Dieterich's (1994) formula: Seismicity rate in response to shear and normal stress changes: $N(t) = \frac{1}{\cdots} \cdot \frac{v_{tec}}{v_{tec}}$



 $\cdot \left\{ \dot{\tau}_{tec} \cdot \left\{ 1 + \frac{p(t)}{\sigma_n^0} \right\} + g_0 \cdot \dot{p}(t) \right\} \right\}$

time in hr

nts





Fig. 5: Comparison between the CPT and RST results for total seismicity of the two-cycle injection with D=0.1 m2/s and lower trigger threshold of 1 kPa after integration over the entire volume of fluid infiltration. There are no visible differences.

Reference

J. Dieterich, A Constitutive Law for the Rate of

$$\frac{d\gamma}{dt} = \frac{1}{A \cdot \sigma_n(t)} \cdot \left\{ 1 - \gamma(t) \cdot \left\{ \frac{\tau(t)}{\sigma_n(t)} - \alpha \right\} \cdot \frac{d\sigma_n}{dt} \right\}$$

background seismicity: v_{tec} natural stress rate: $\dot{\tau}_{tec}$ If normal stress changes due to fluid injection:

Or written as Ricatti equation

$$N(t) = R(t) \cdot v_{tec}$$

 $g_0 = \frac{\tau_{tec}}{\sigma^0} - \alpha$

 $\sigma_n(t) = \sigma_n^0 - p(t)$

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$$N(t) = D(t) \cdot \alpha$$