

Ground Motion Prediction for Induced Seismicity Using the Ambient Seismic Field

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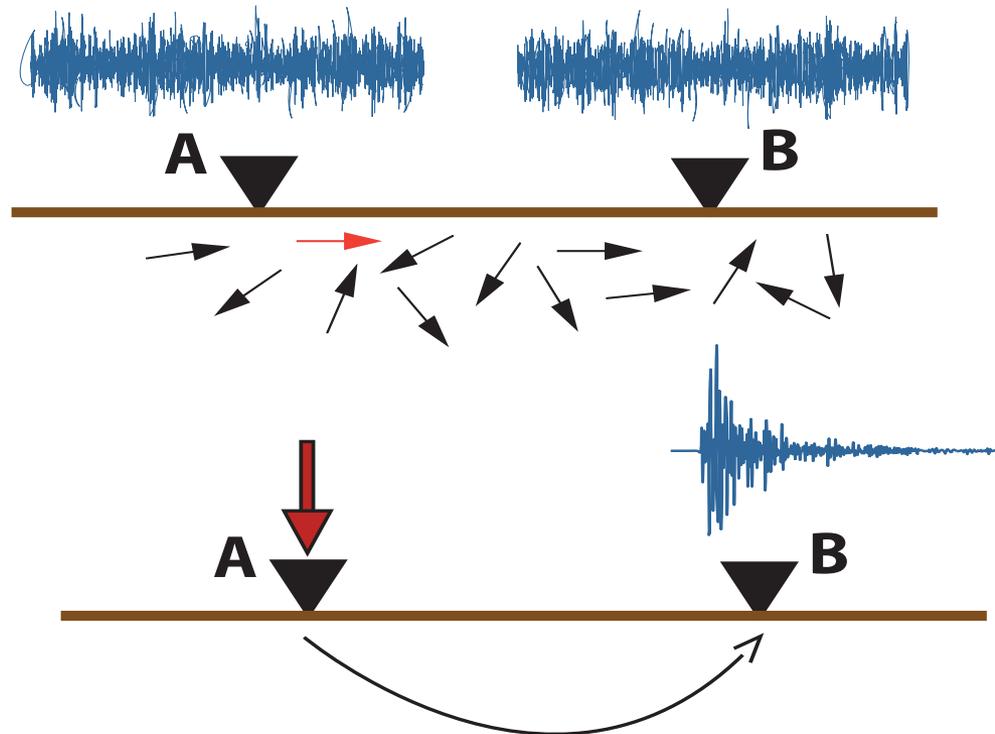
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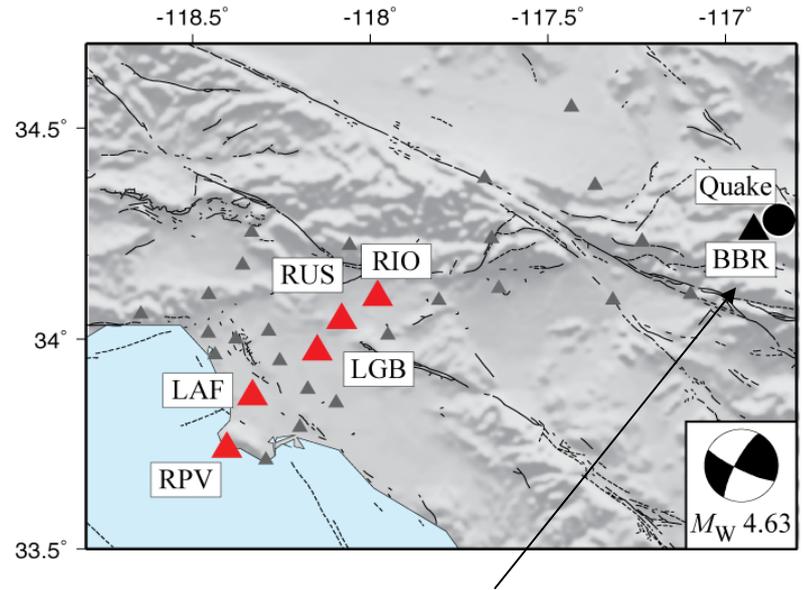
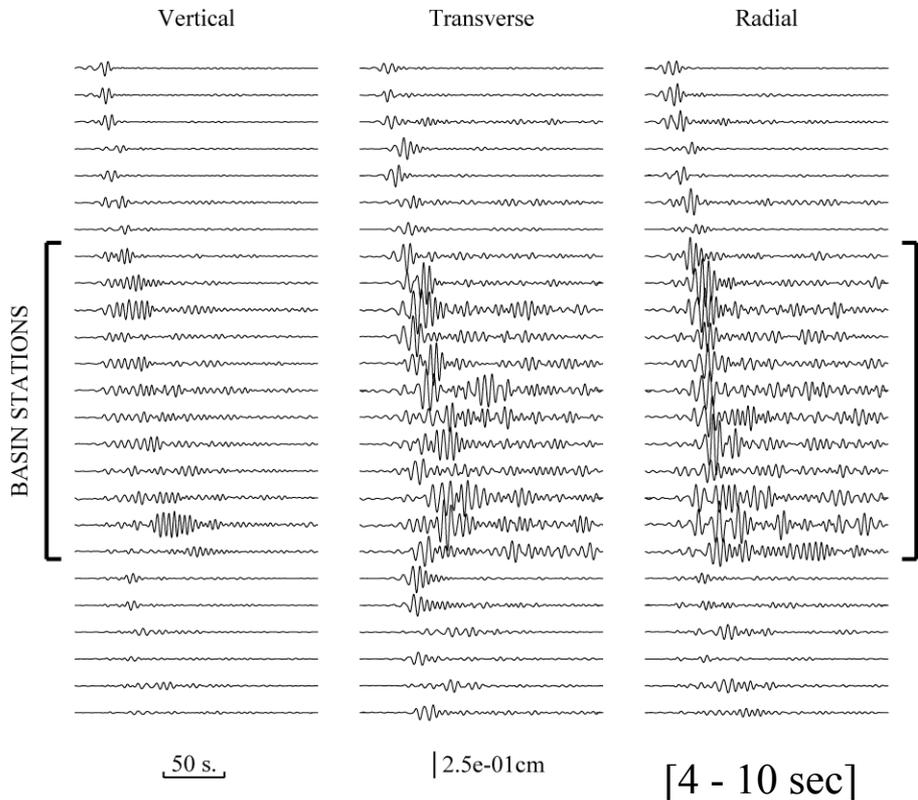
Ambient-Field Green's Functions



Cross-correlation of ground motion at two stations converges to the impulse response with sufficient averaging

Proof of Concept for M 4.6 Big Bear EQ

Earthquake records

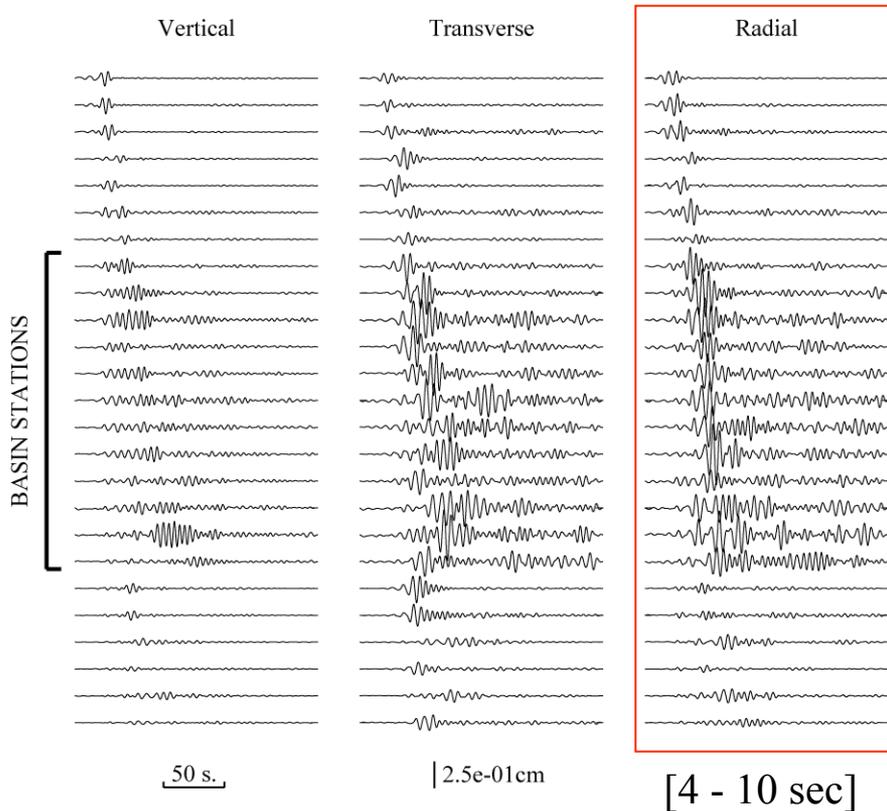


10 Feb 2001 M_w 4.6 Big Bear
Graves (2008)

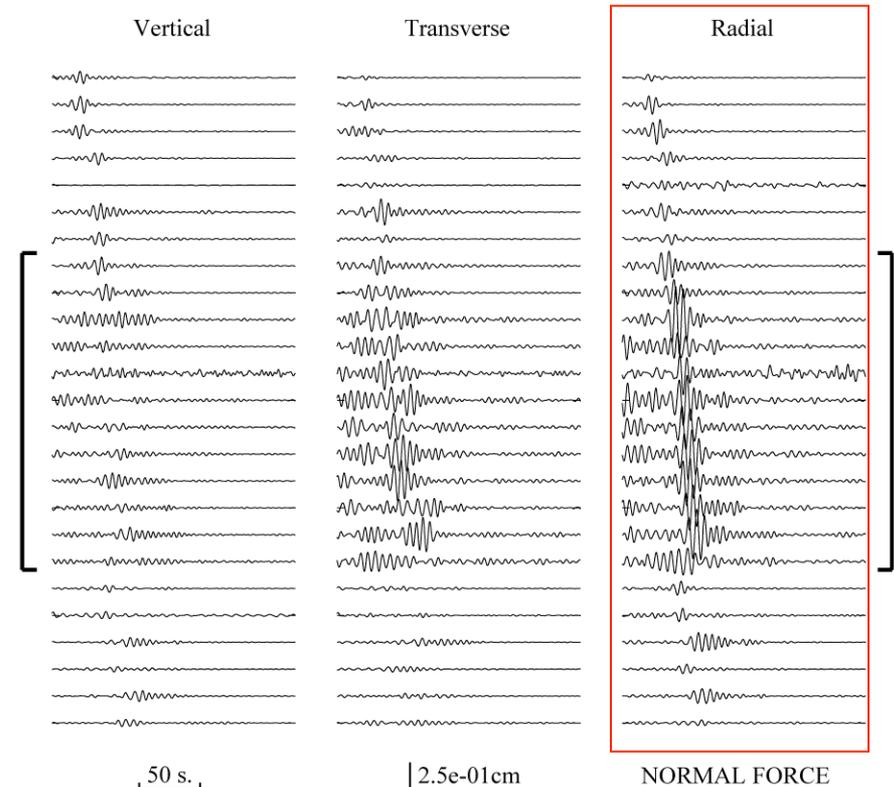
Amplitudes, duration, and waveform complexity are greater in the basin

Proof of Concept for *M* 4.6 Big Bear EQ

Earthquake records



Impulse response records



Prieto and Beroza (2008)

Amplitudes, duration, and waveform complexity in the basin are reproduced by the ambient field.

Approaches to Ground Motion Prediction

Data-Driven

Model-Guided

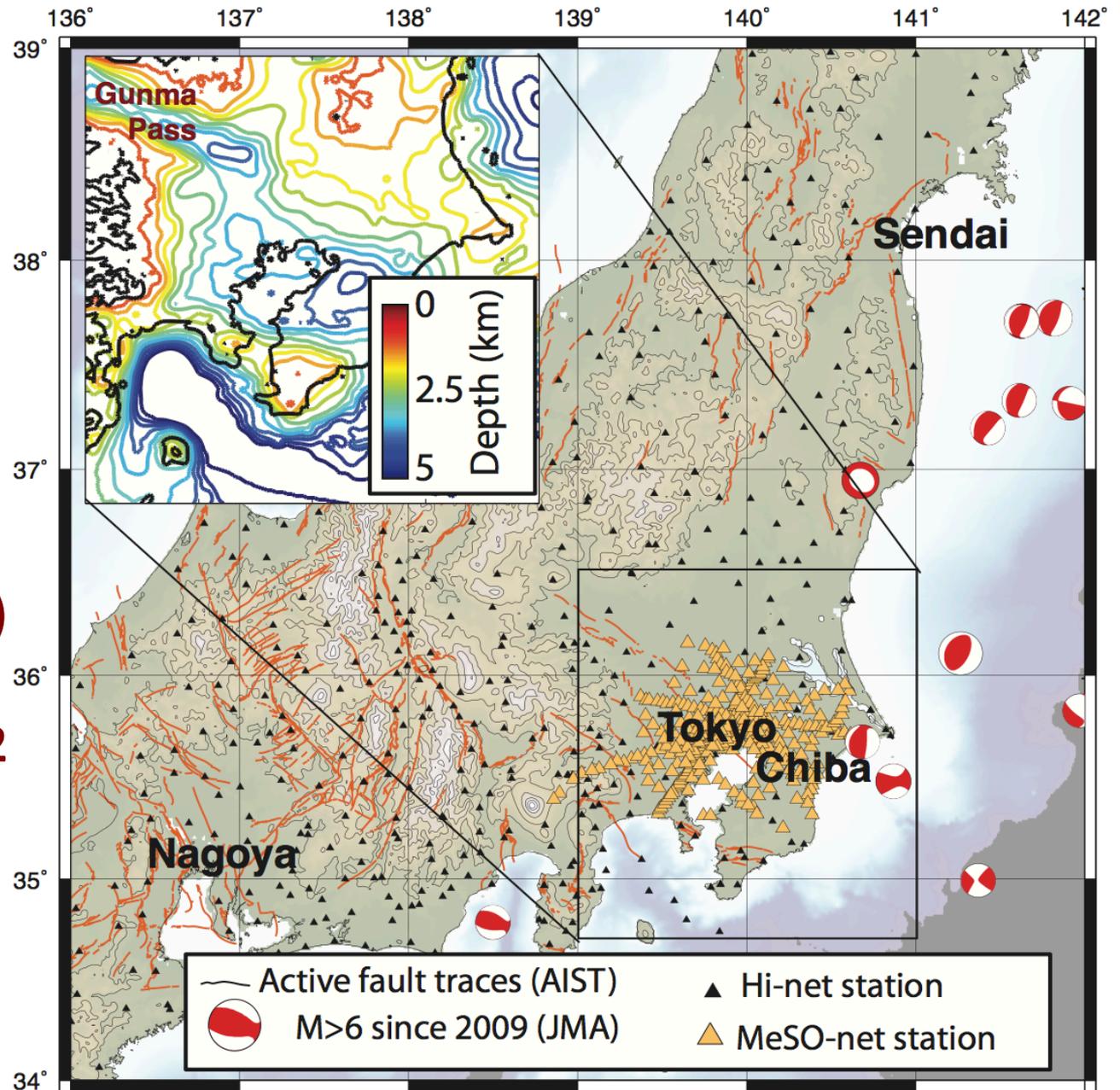


GMPEs

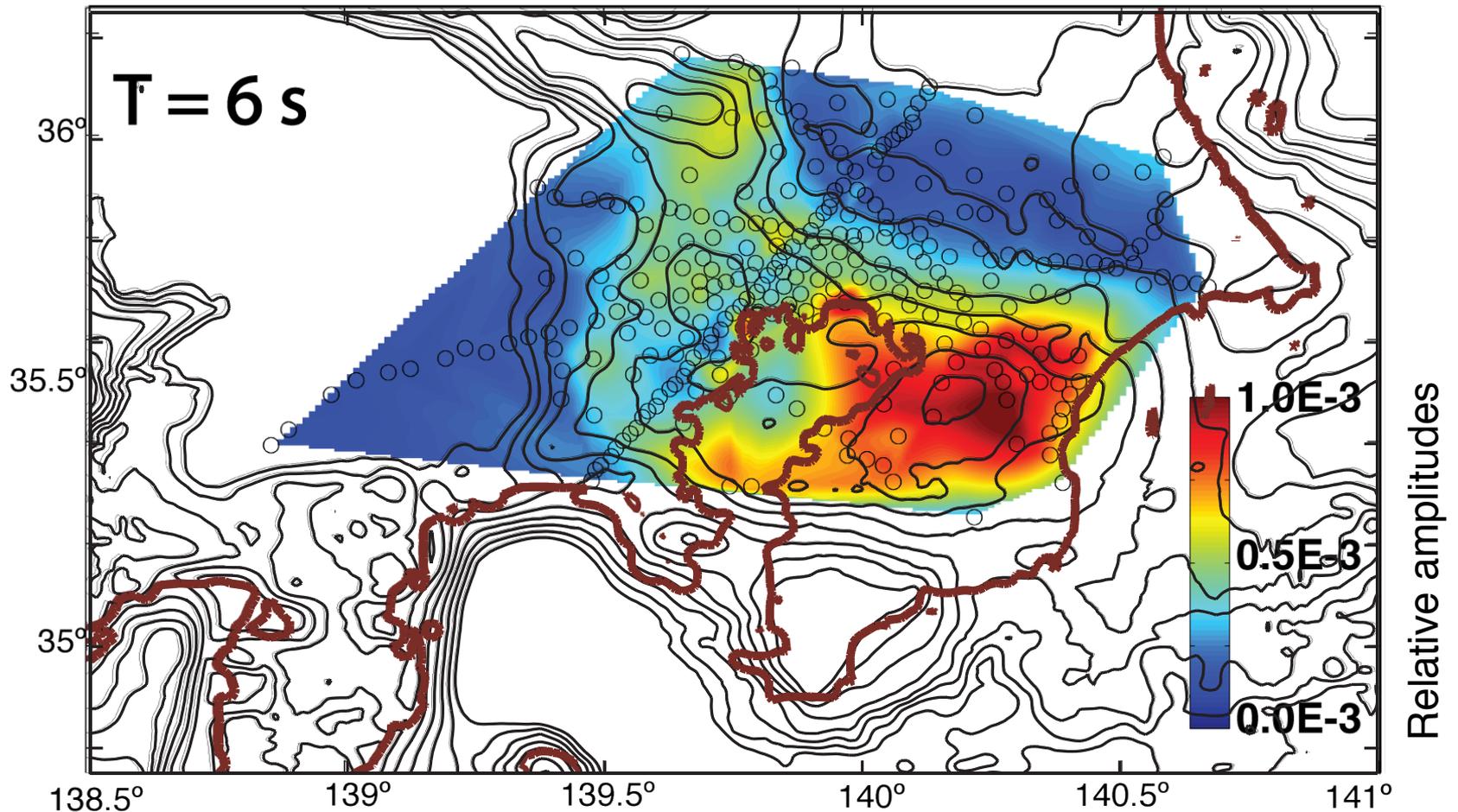
Fit ground motions
to parameterized
equations.

How can ambient-
field measurements
contribute?

Hi-Net
(deep borehole)
+
MESO-Net
(shallow borehole)
 $\#GFs = 9 \times N(N-1)/2$



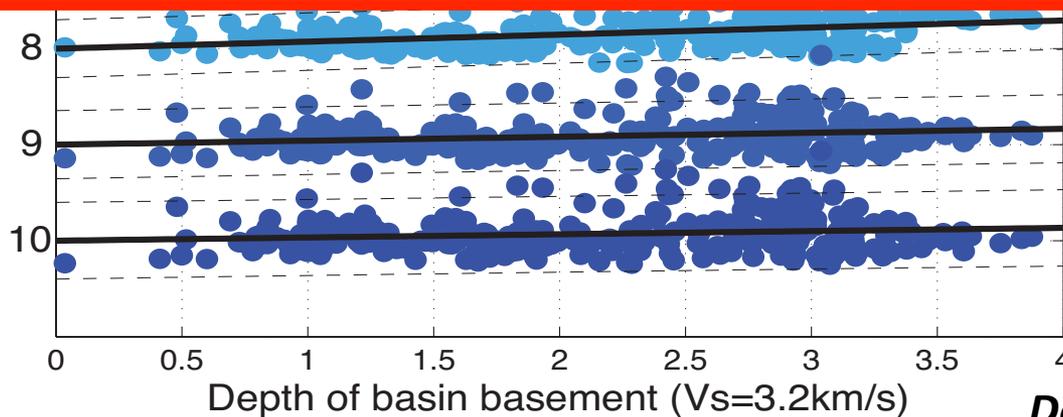
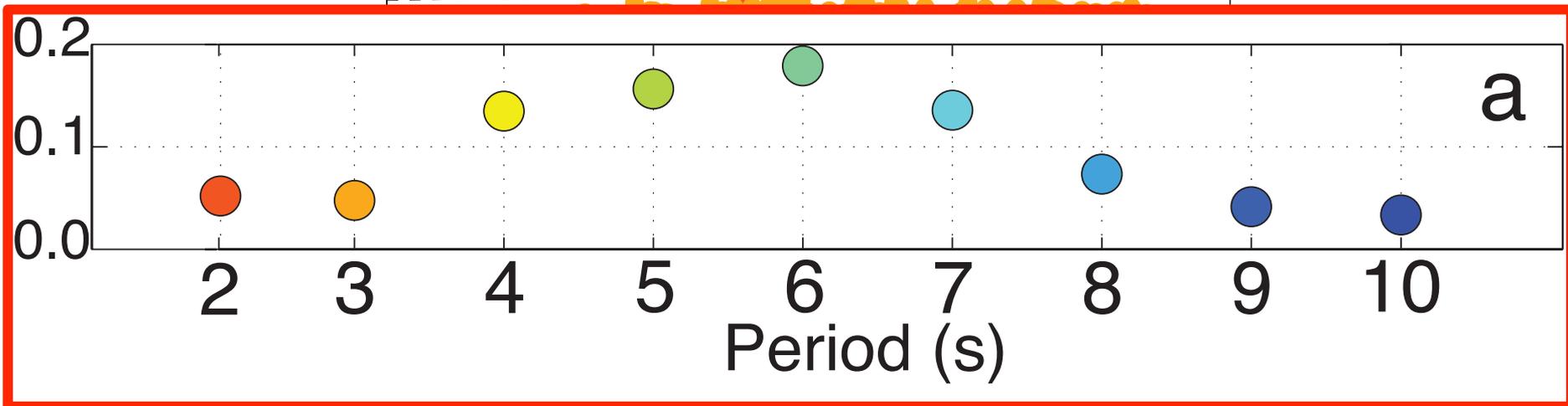
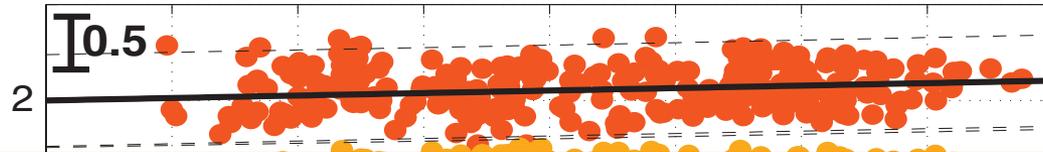
Average Basin Amplification vs. Basin Depth



Contours show basin depth – Colors show amplification

Basin Amplification ~ Basin Depth

$$\log A = a Z + b$$



Approaches to Ground Motion Prediction

Data-Driven

Model-Guided



GMPEs

Simulations

Fit ground motions
to parameterized
equations.

Model wave
propagation through
imaged 3D structure.

How can ambient-
field measurements
contribute?

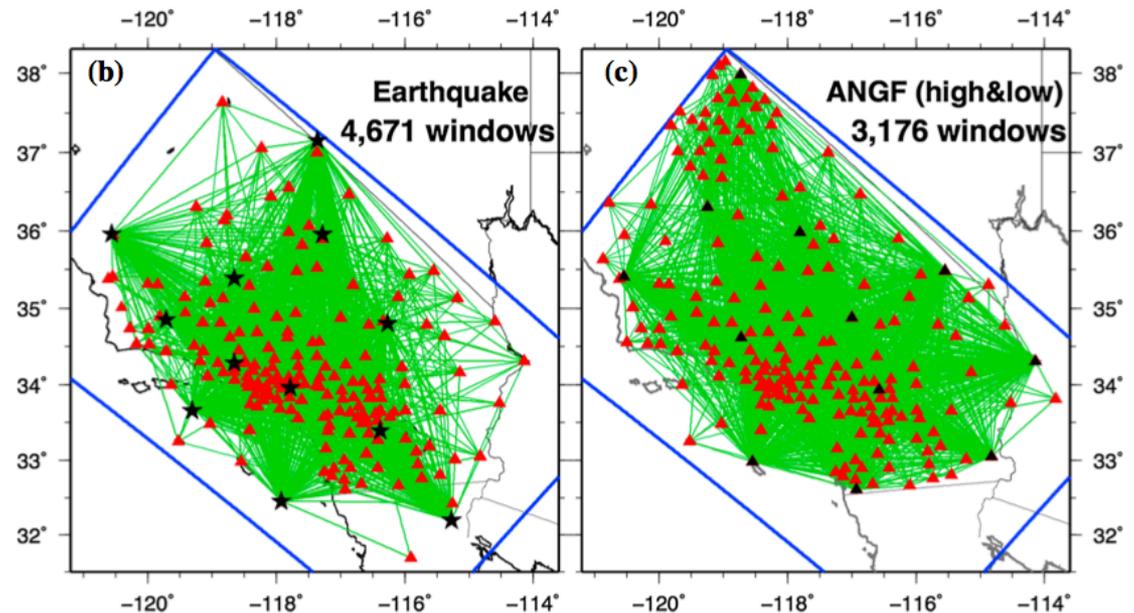
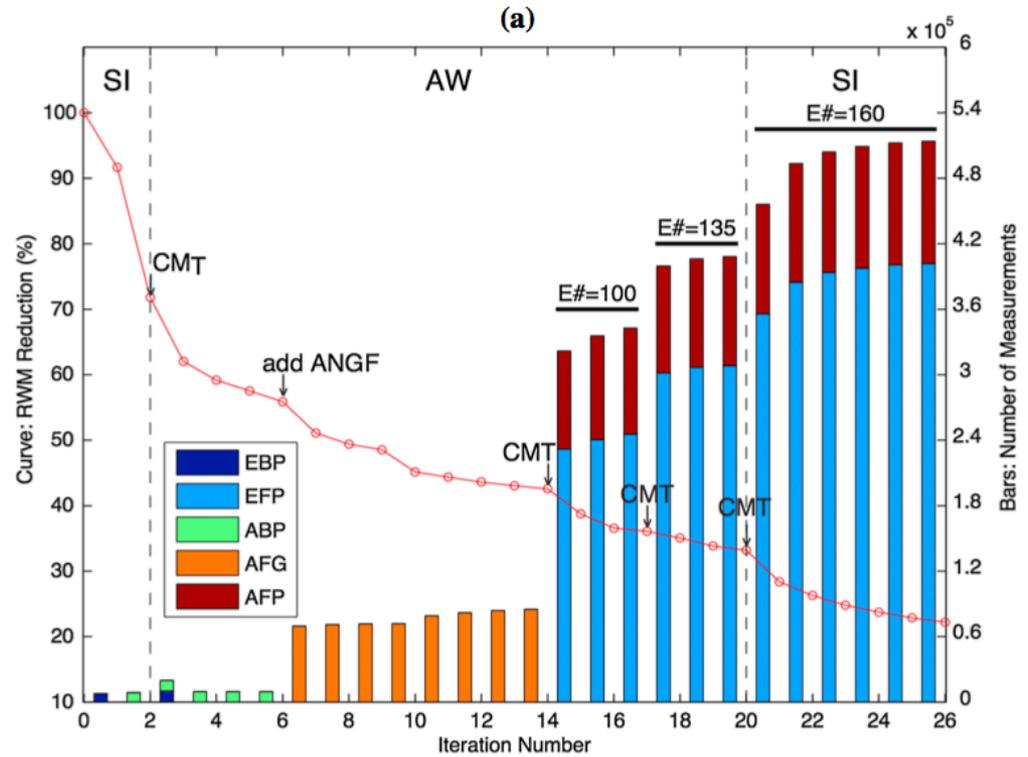


Ambient-Field Observations

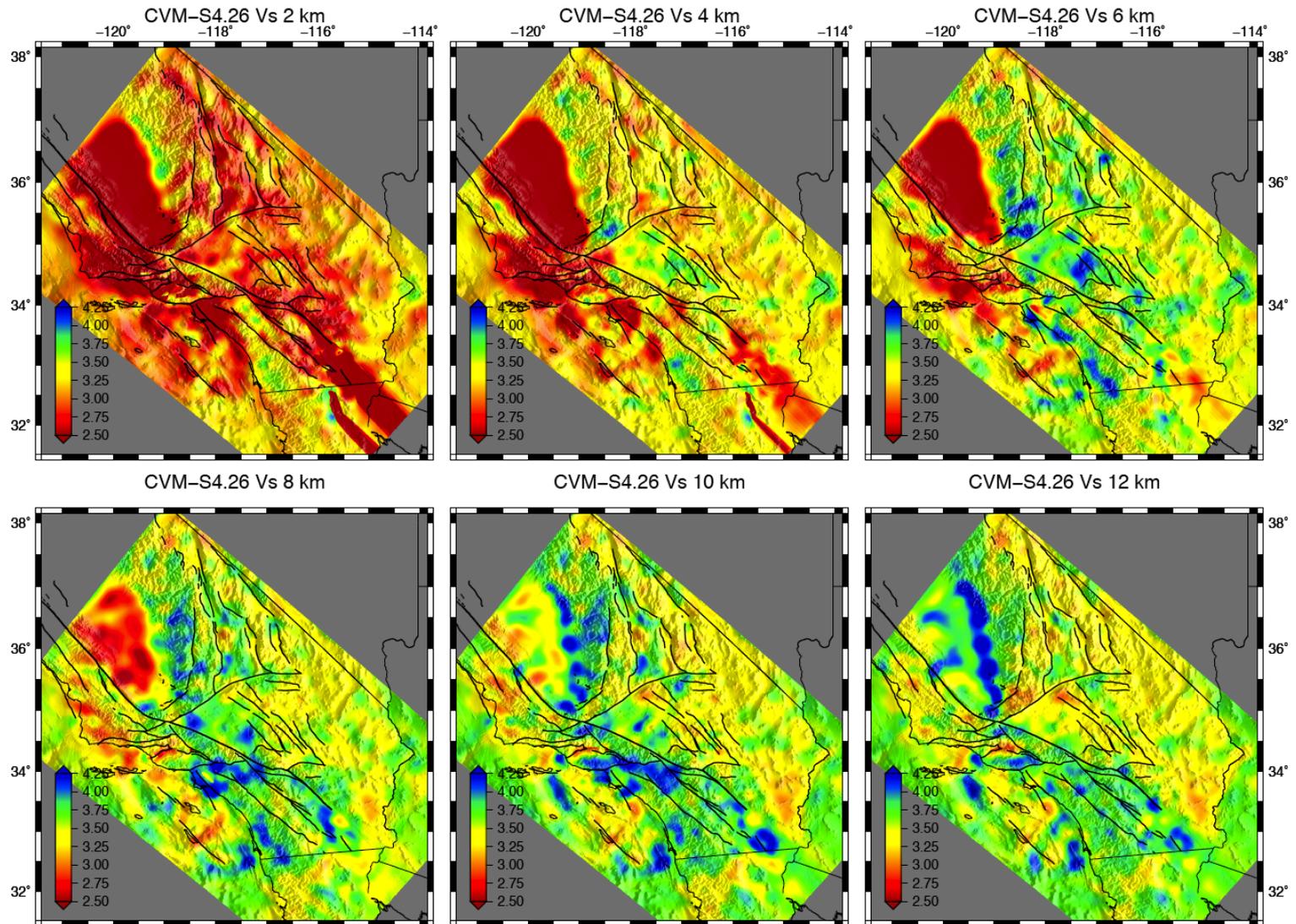
Data even in the absence of earthquakes

Constrain shallow structure relevant to ground motion prediction.

(Lee et al., 2014)

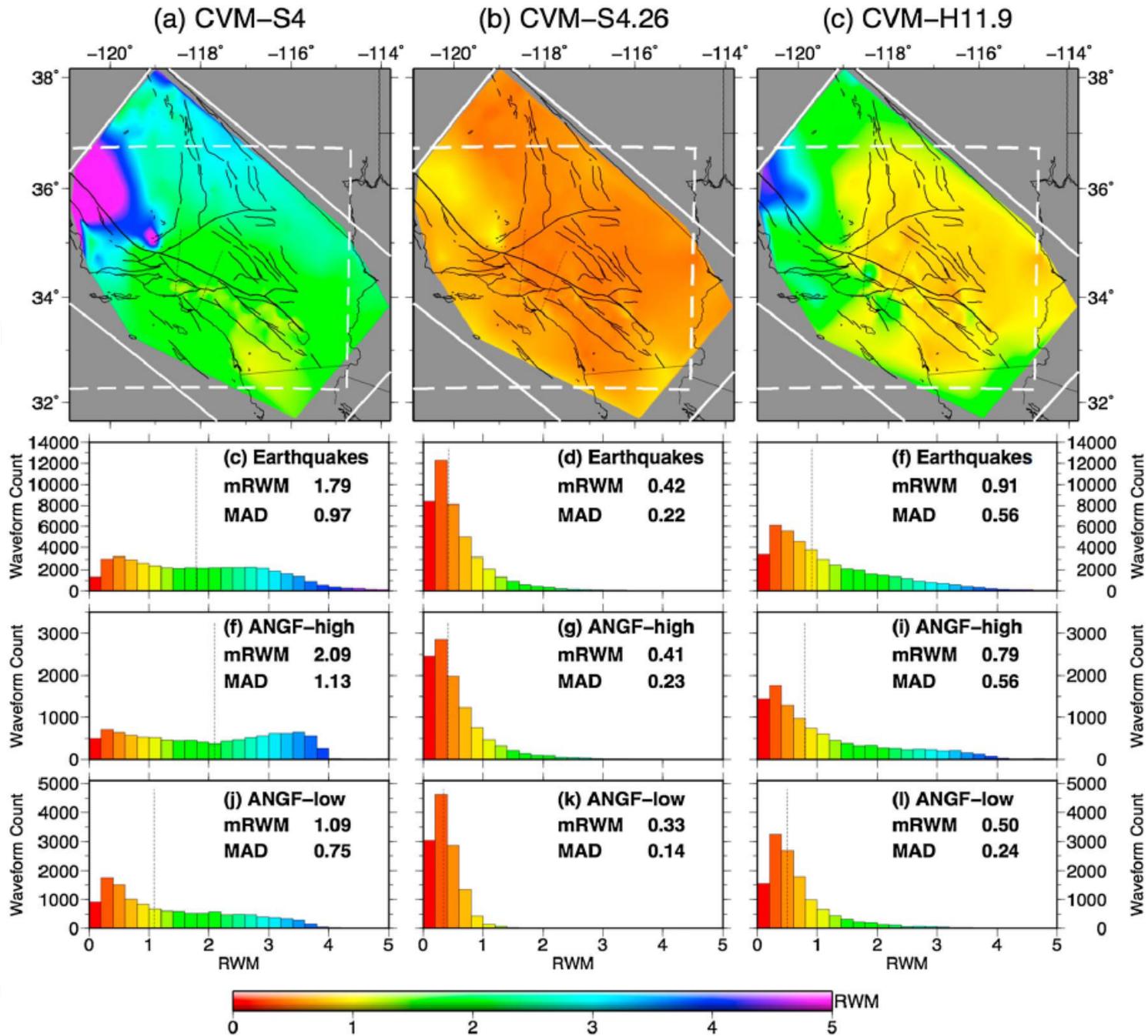


Full 3D Tomography of Southern California



(Lee et al., 2014)

Reduced Waveform Misfit



(Lee et al., 2014)

Approaches to Ground Motion Prediction

Data-Driven

Hybrid

Model-Guided



GMPEs

Virtual EQs

Simulations

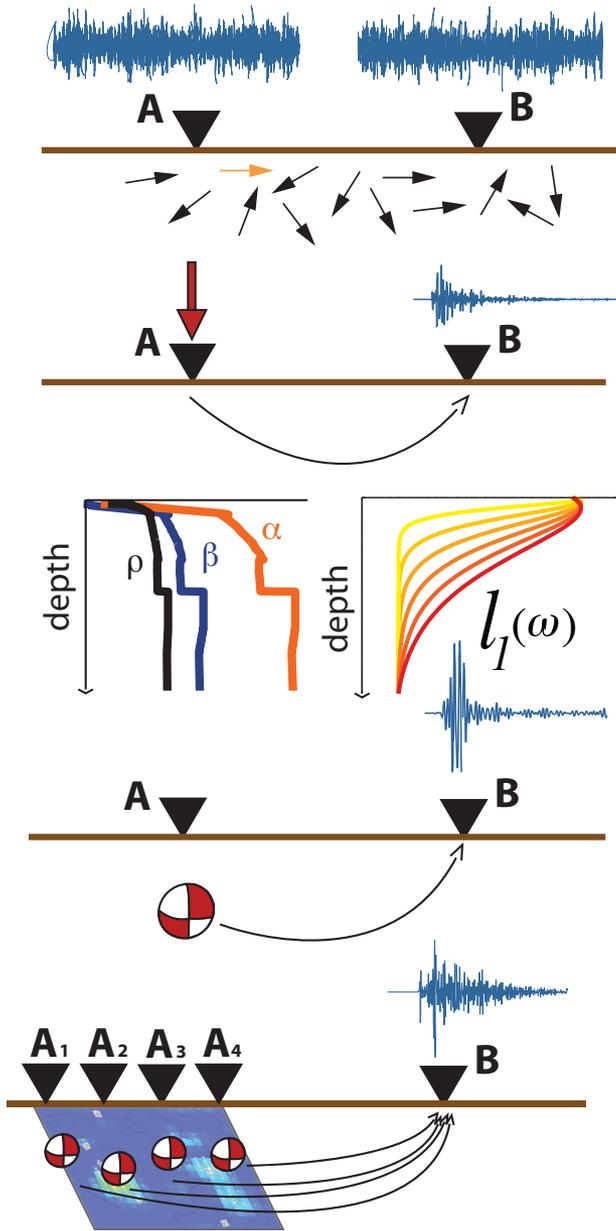
Fit ground motions to parameterized equations.

Use ambient seismic field to characterize effect of 3D structure.

Model wave propagation through imaged 3D structure.



Virtual Earthquake Method



Continuous Recording of ASF



Earth Response to an Impulsive Force



Source-Depth Correction



Double-Couple Correction



Finite-Source Correction

Denolle et al. (2013)

Depth and Mechanism Corrections

Love:

$$\hat{U}_T(\vec{\xi}, \omega) = \frac{1}{l_1(\vec{\xi}', 0, \omega)} \left[-ik_L(\omega) \hat{M}_{TR}(\omega) l_1(\vec{\xi}', h, \omega) + \hat{M}_{TD}(\omega) l_1'(\vec{\xi}', h, \omega) \right] \hat{G}_{TT}(\vec{\xi}', \vec{\xi}, \omega)$$

Rayleigh:

$$\begin{aligned} \hat{U}_R(\vec{\xi}, \omega) &= \frac{1}{r_1(\vec{\xi}', 0, \omega)} \left[-ik_R \hat{M}_{RR}(\omega) r_1(\vec{\xi}', h, \omega) + \hat{M}_{RD}(\omega) l_1'(\vec{\xi}', h, \omega) \right] \hat{G}_{RR}(\vec{\xi}', \vec{\xi}, \omega) + \\ &\quad \frac{1}{r_2(\vec{\xi}', 0, \omega)} \left[-ik_R \hat{M}_{DR}(\omega) r_1(\vec{\xi}', h, \omega) + \hat{M}_{DD}(\omega) l_1'(\vec{\xi}', h, \omega) \right] \hat{G}_{DR}(\vec{\xi}', \vec{\xi}, \omega), \\ \hat{U}_D(\vec{\xi}, \omega) &= \frac{1}{r_1(\vec{\xi}', 0, \omega)} \left[-ik_R \hat{M}_{RR}(\omega) r_1(\vec{\xi}', h, \omega) + \hat{M}_{RD}(\omega) l_1'(\vec{\xi}', h, \omega) \right] \hat{G}_{RD}(\vec{\xi}', \vec{\xi}, \omega) + \\ &\quad \frac{1}{r_2(\vec{\xi}', 0, \omega)} \left[-ik_R \hat{M}_{DR}(\omega) r_1(\vec{\xi}', h, \omega) + \hat{M}_{DD}(\omega) l_1'(\vec{\xi}', h, \omega) \right] \hat{G}_{DD}(\vec{\xi}', \vec{\xi}, \omega), \end{aligned}$$

Make 1D approximation in excitation

All 9 Components of GFs (2-10 s)

P-SV

Z

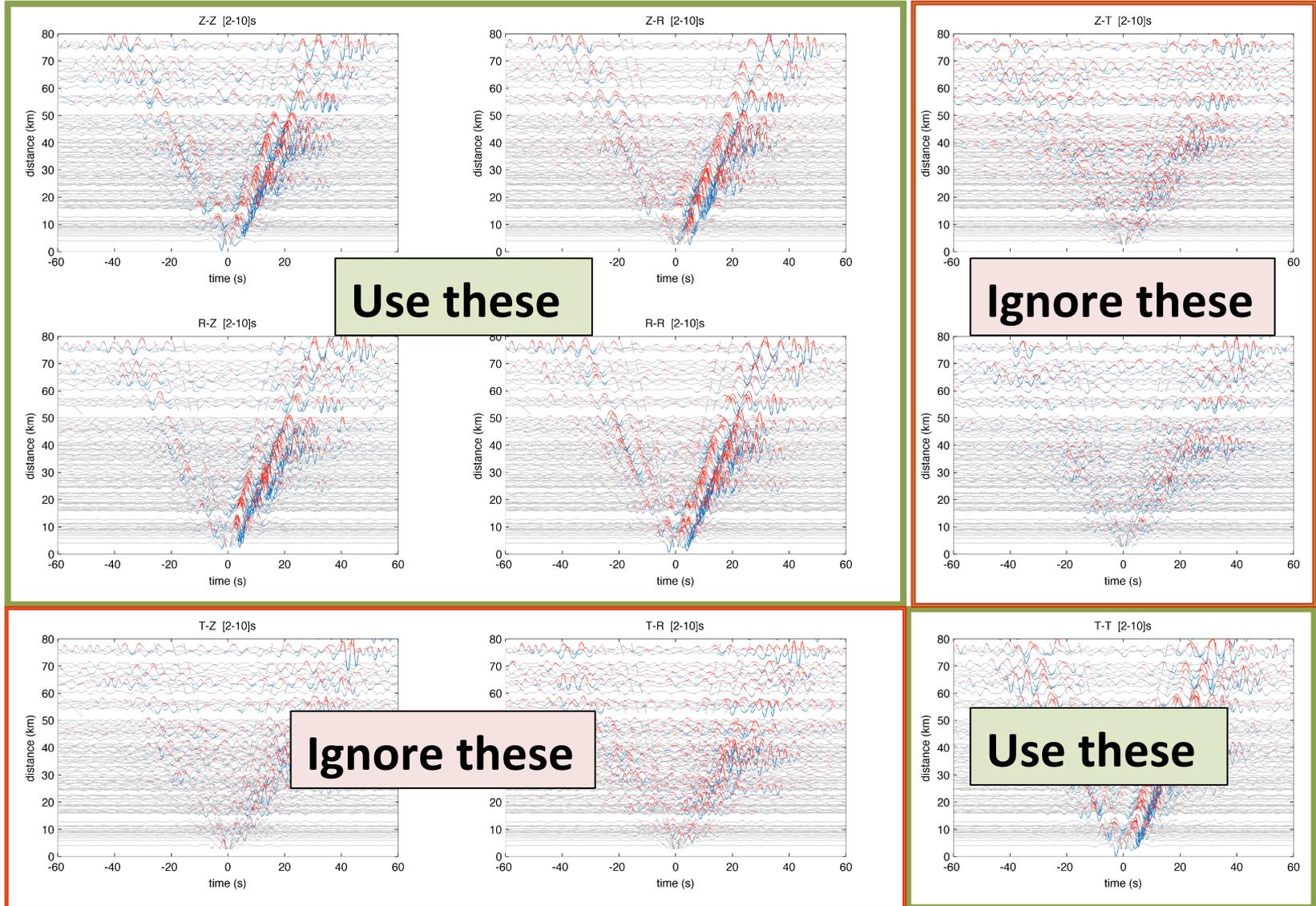
R

T

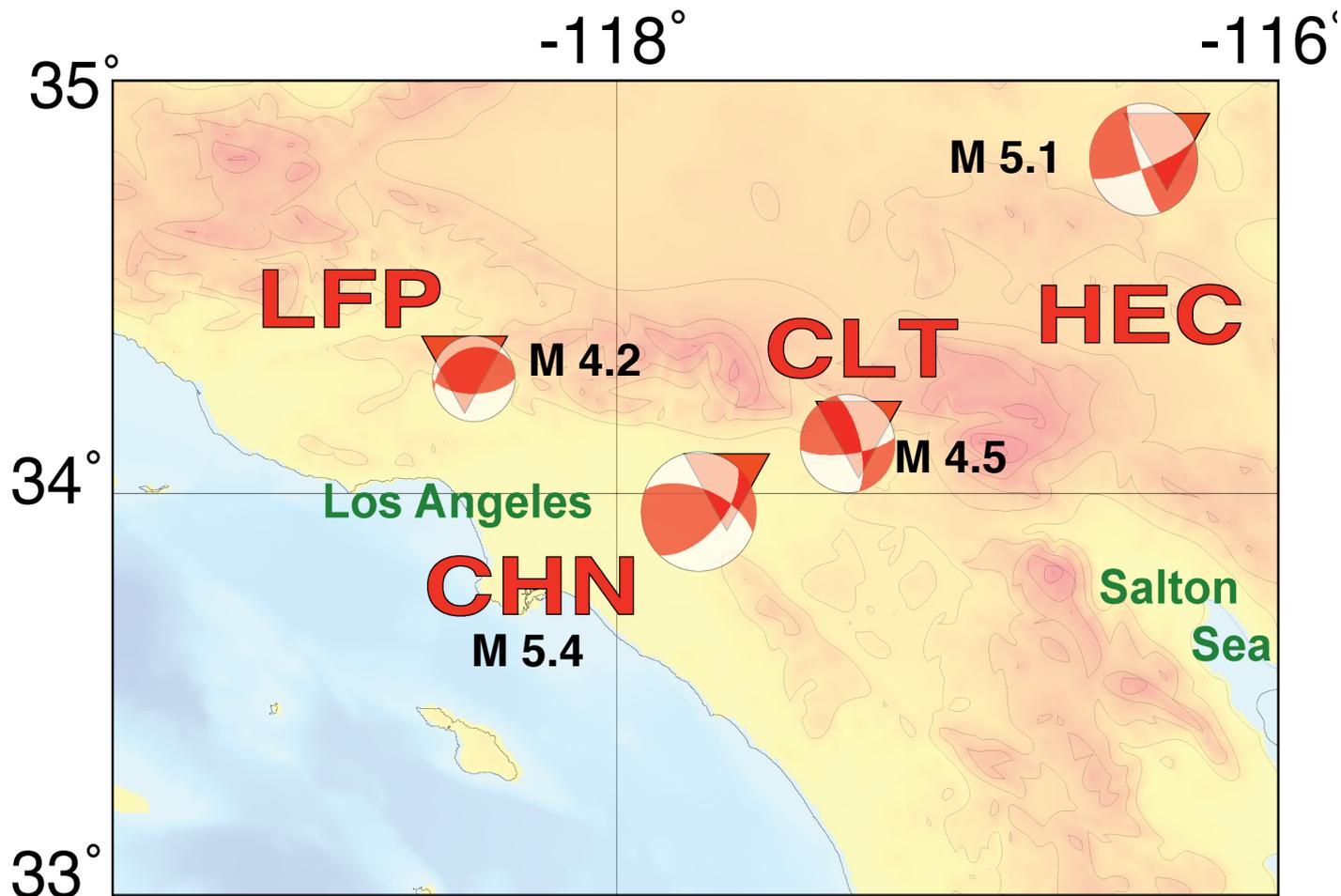
Z

R

T



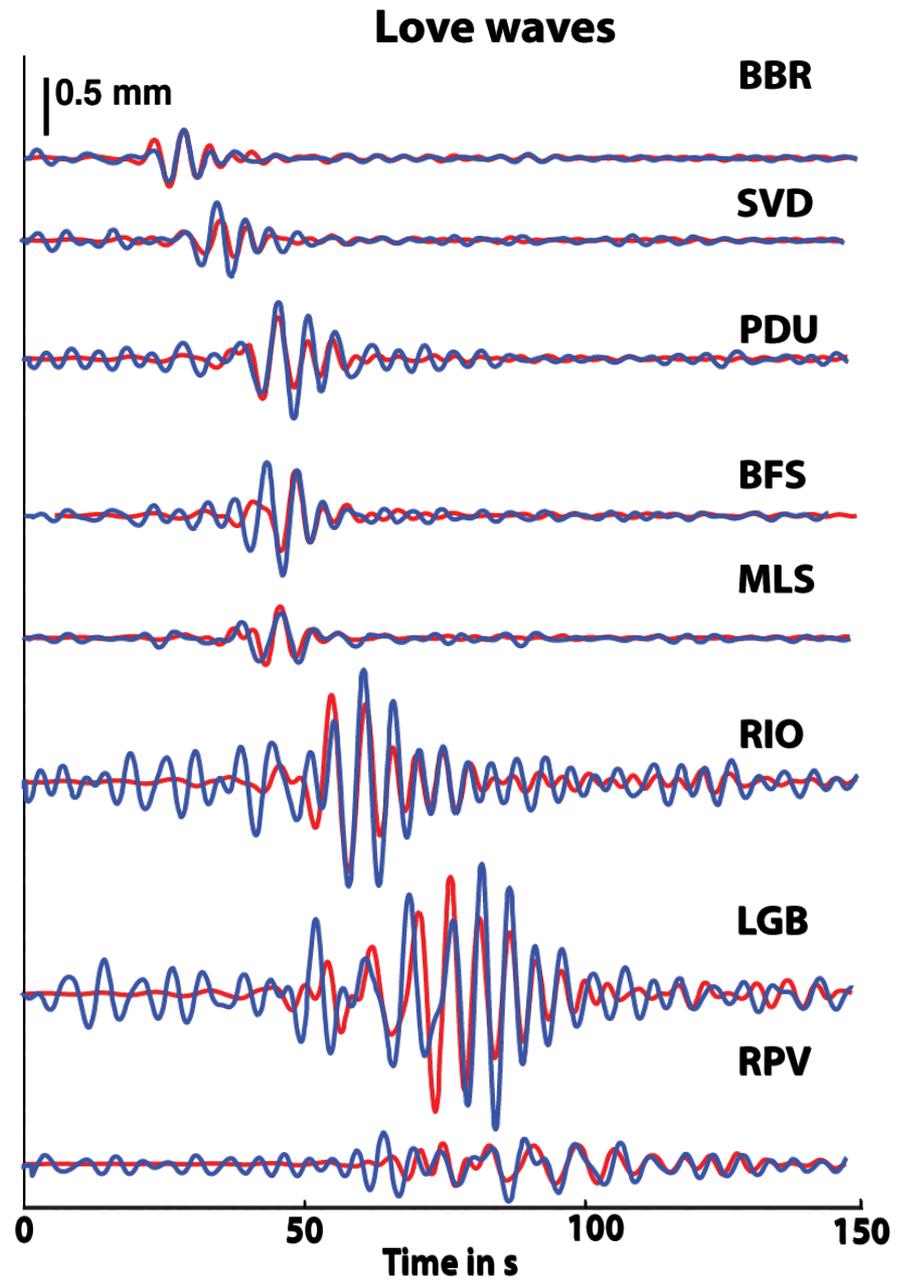
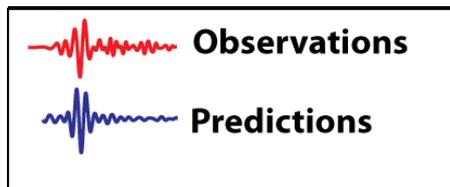
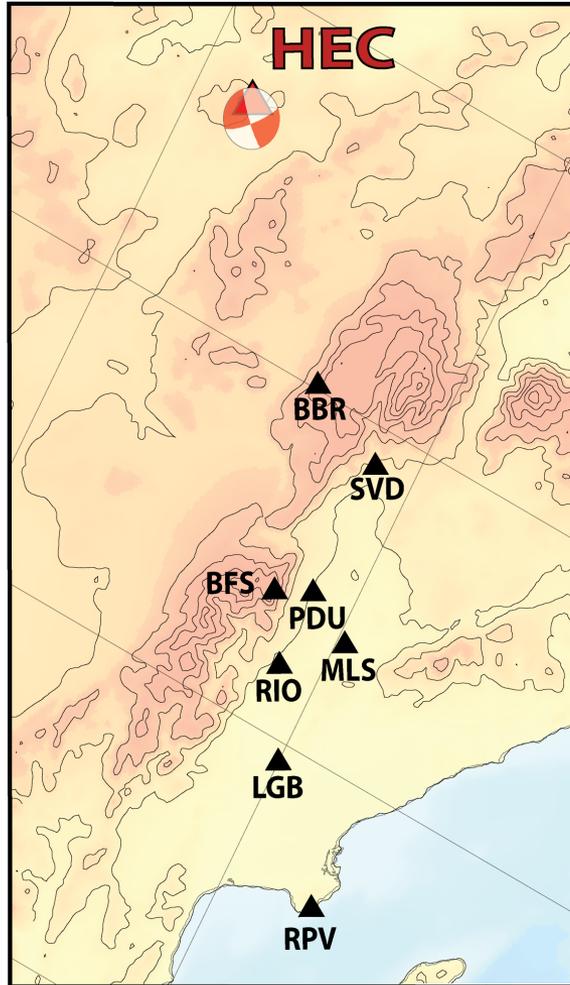
Validate Against 4 Moderate Earthquakes



Each located near a BB station for which we construct a virtual earthquake to compare with seismograms from real earthquake

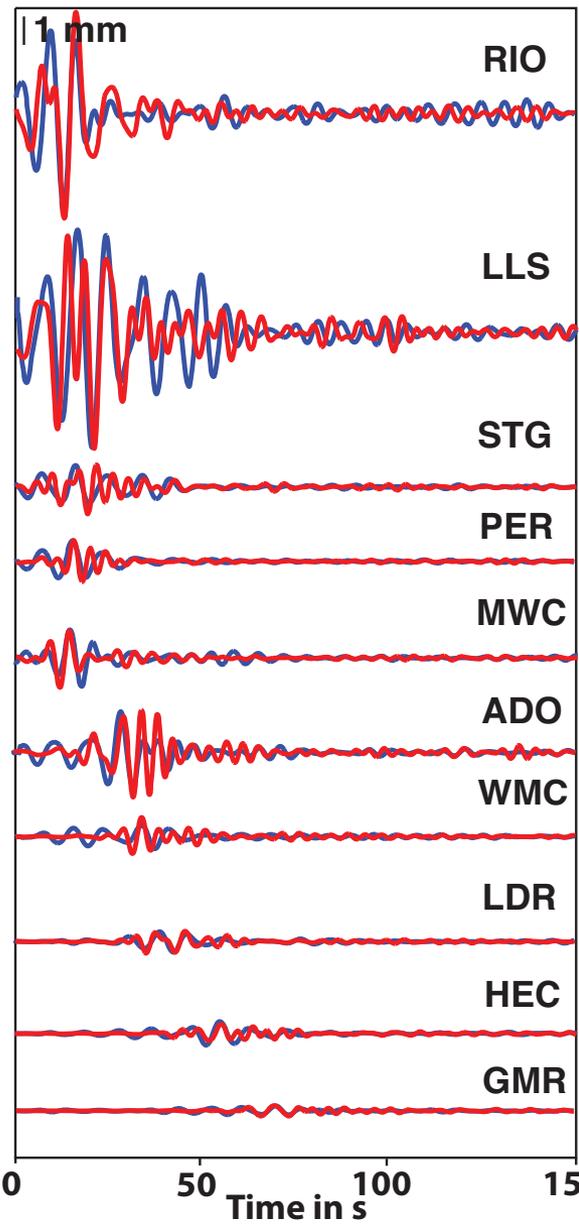
Denolle et al. (2013)

Hector Road M 5.1

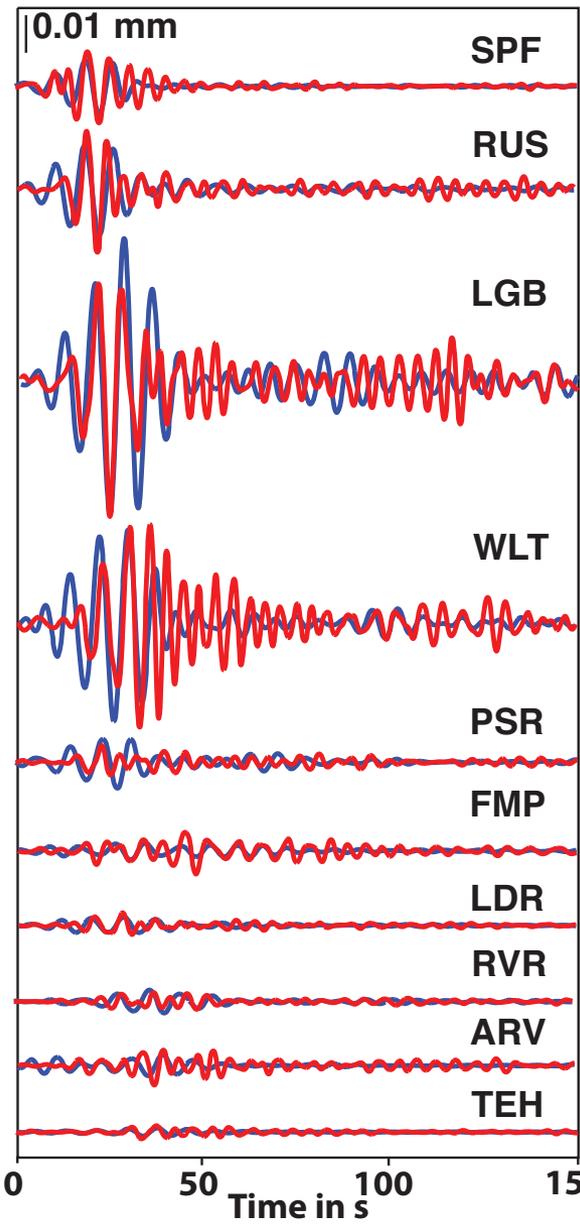


Further Validation with Moderate Earthquakes

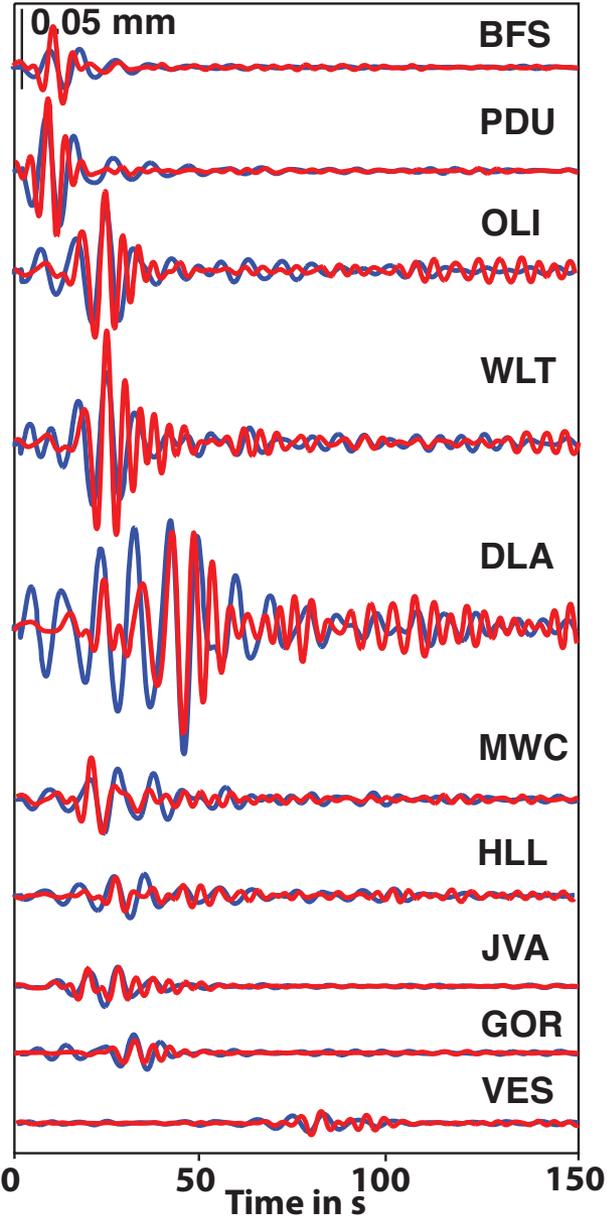
Chino Hills



San Fernando



San Bernardino



Application to GMP for Induced Seismicity

Linearity (weak motion)

Shallow - depth correction smaller

Amplitude bias – azimuthal/among components

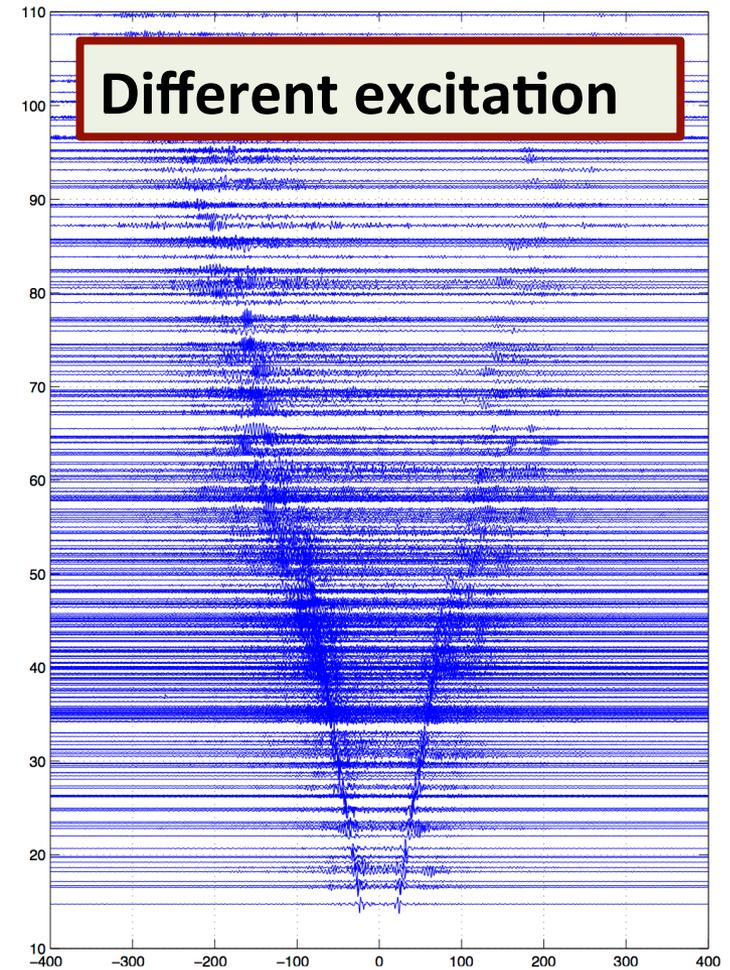
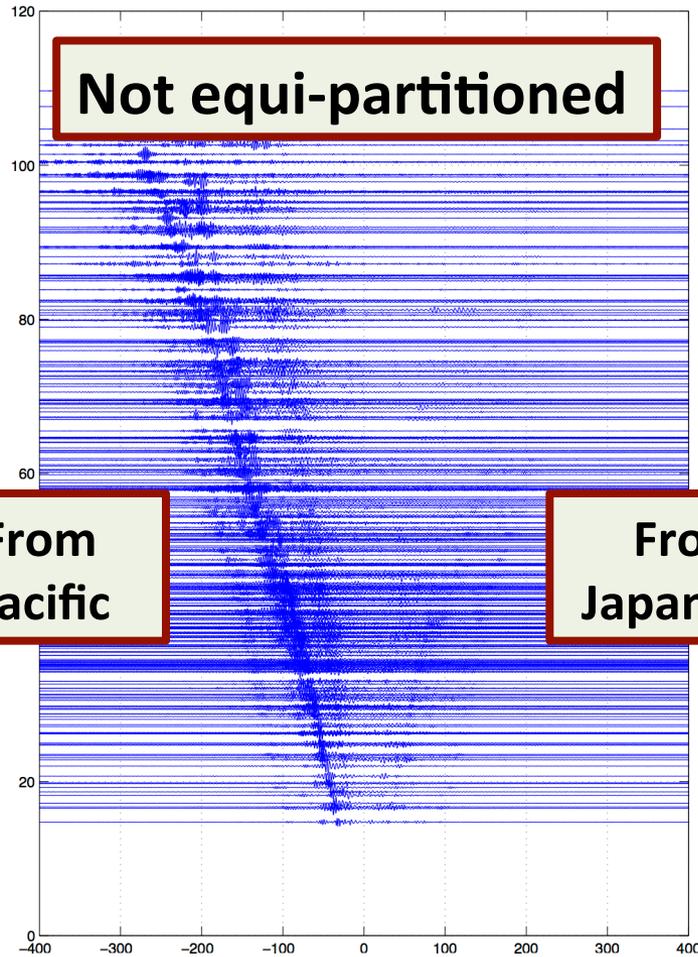
Bandwidth limitation – depends on distance, site

Body waves vs. surface waves

Amplitude Biases – Data from Japan

ZZ

TT

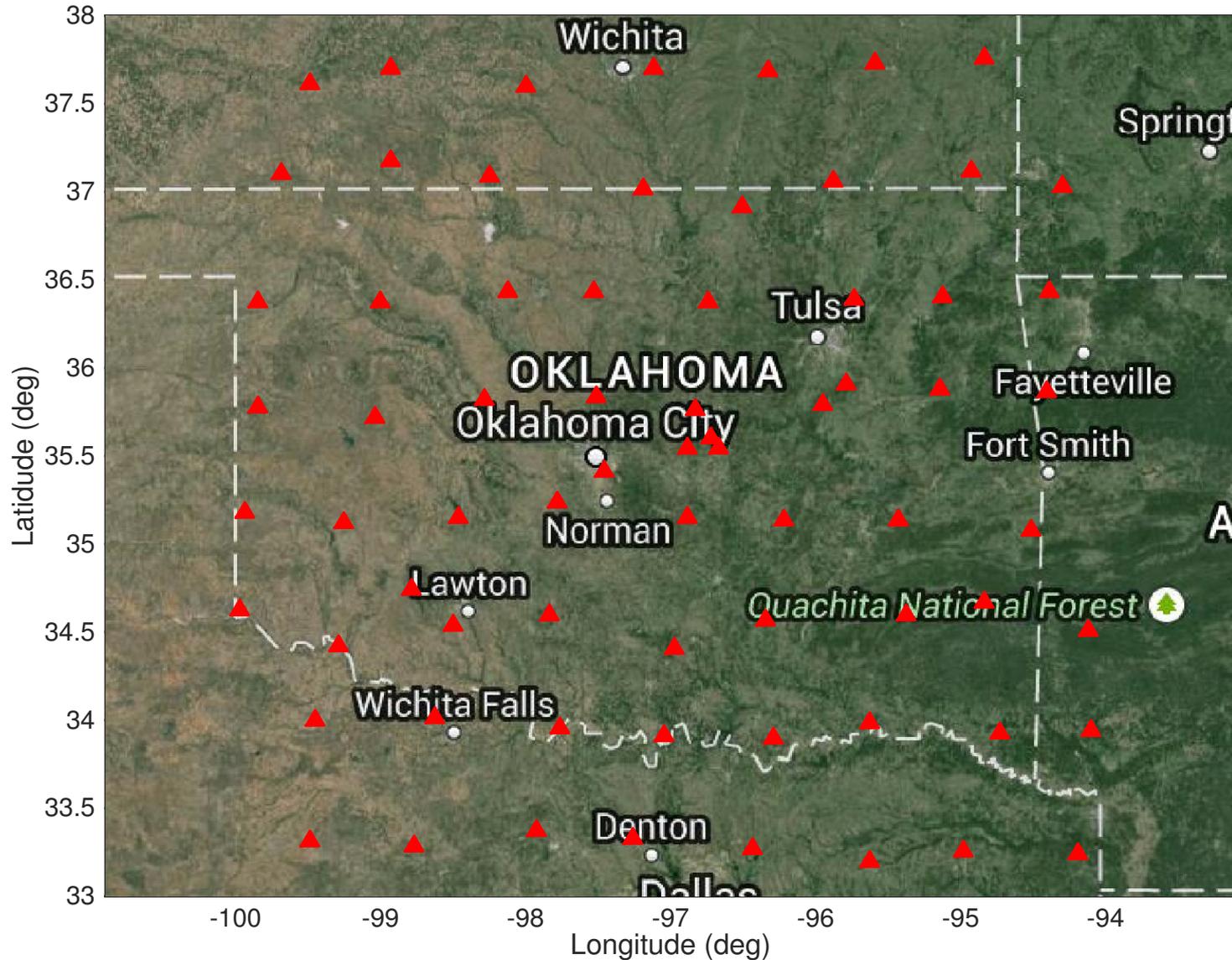


By Azimuth

By Component

Denolle et al. (2014b)

TA Across Oklahoma

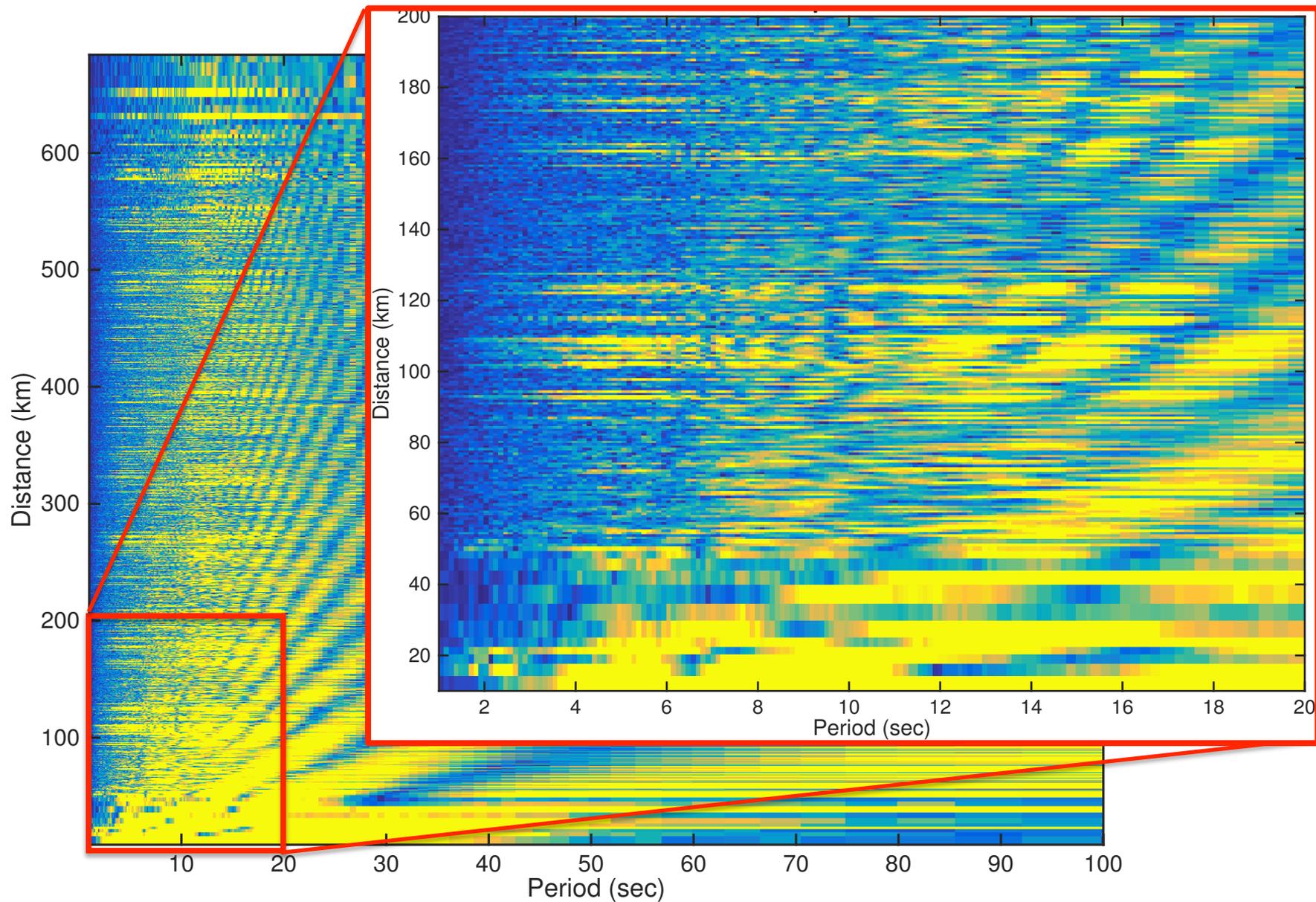


**~1.5 years
continuous
broadband
data**

**~75 km
station
spacing**

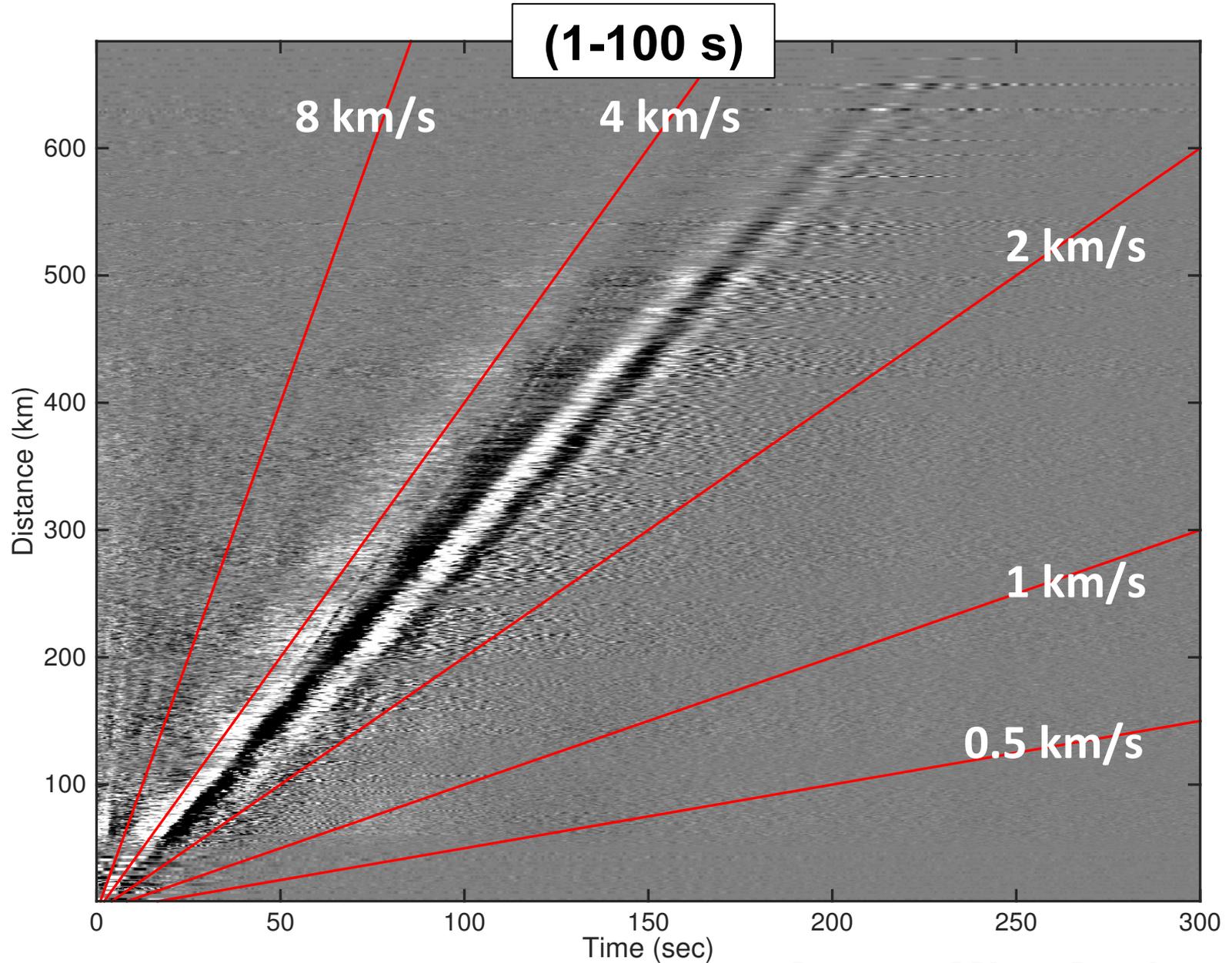
Courtesy of Pierre Boue'

TA Station Pairs - GF Recovery



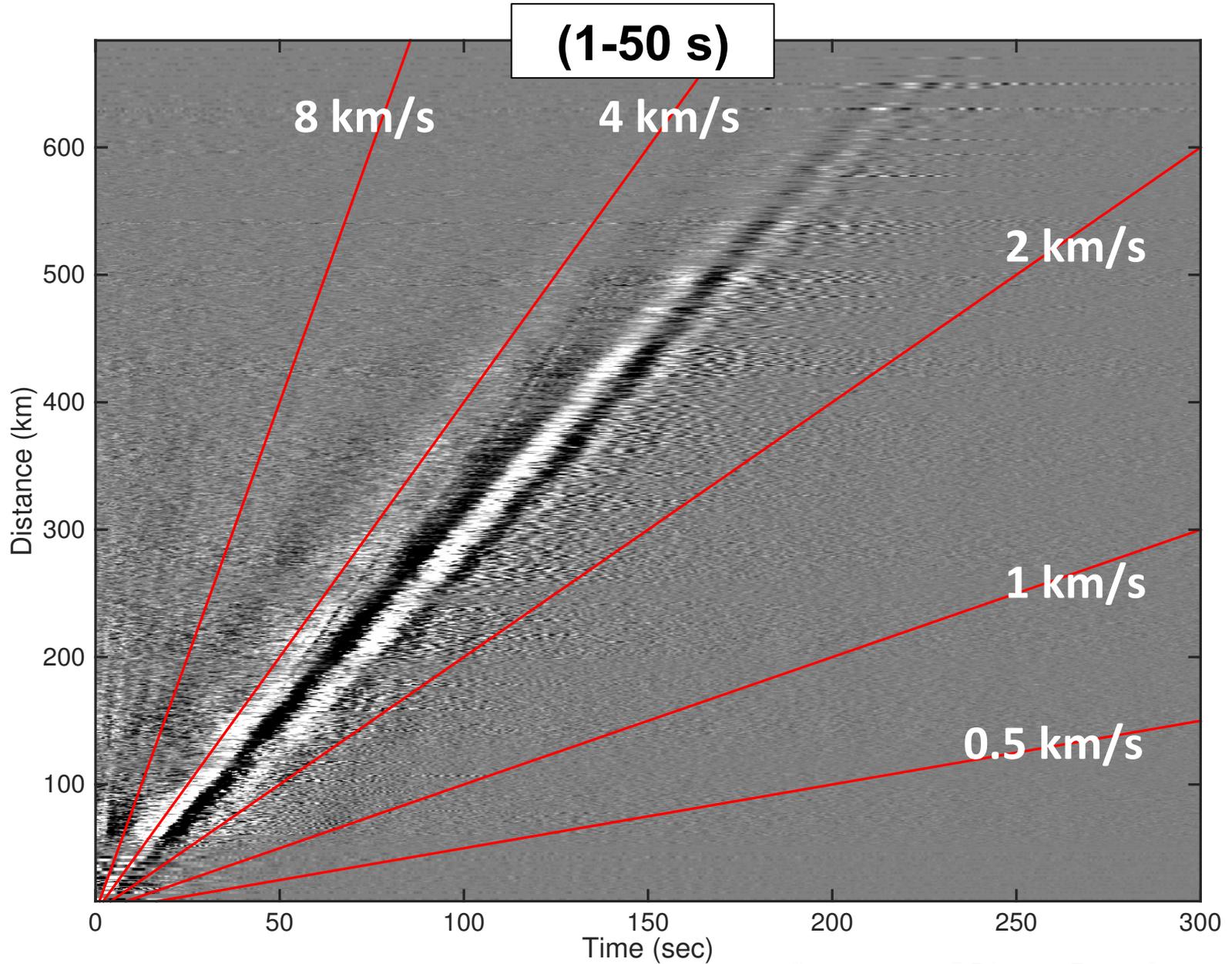
Courtesy of Pierre Boue'

TA Green's Functions



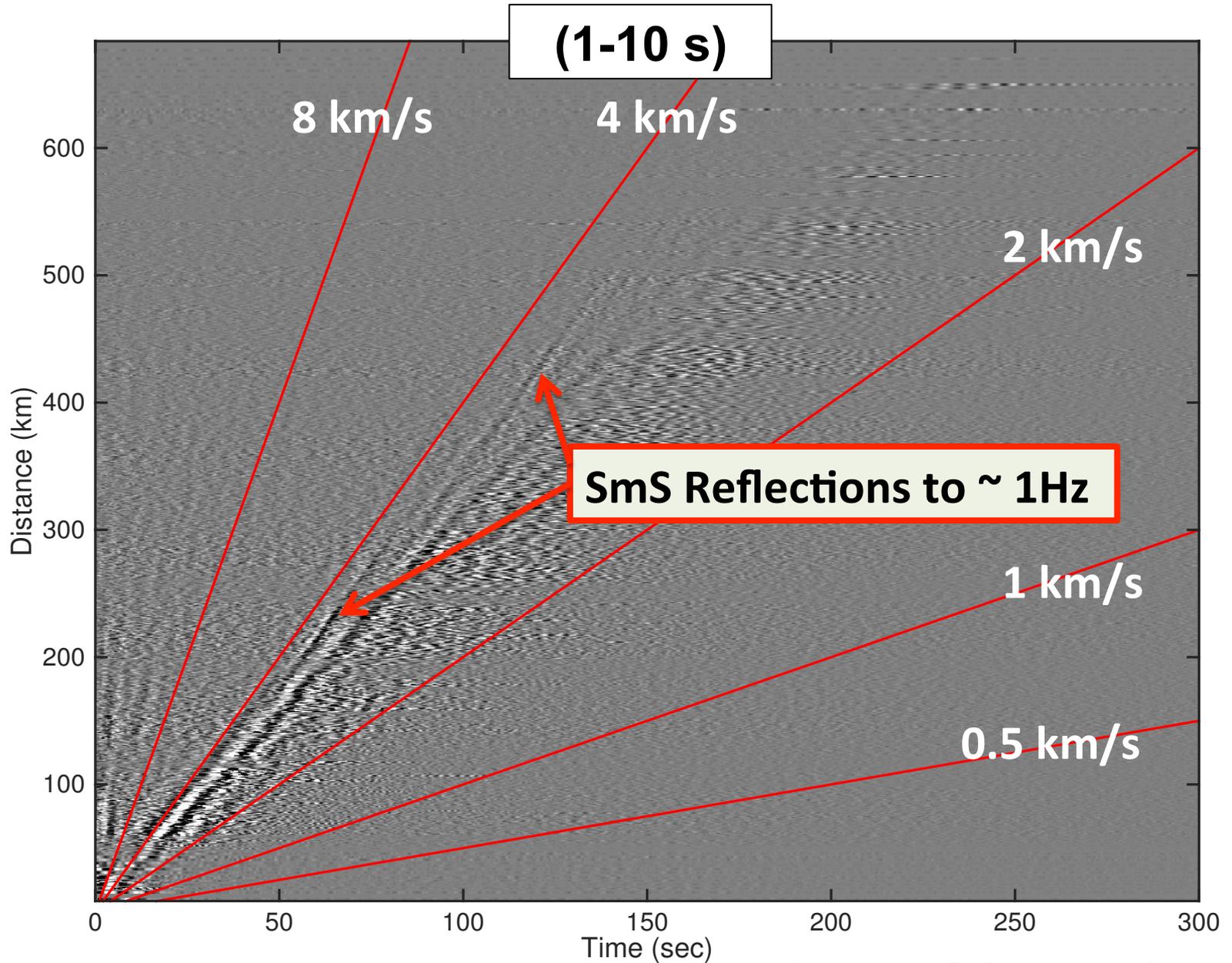
Courtesy of Pierre Boue'

TA Green's Functions



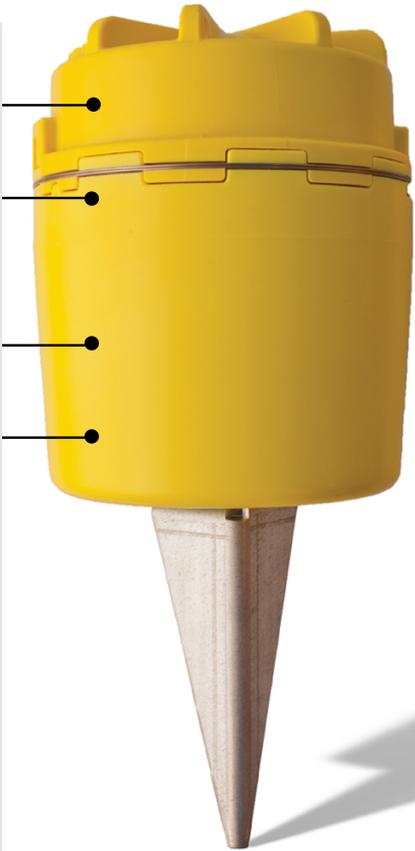
Courtesy of Pierre Boue'

TA Green's Functions



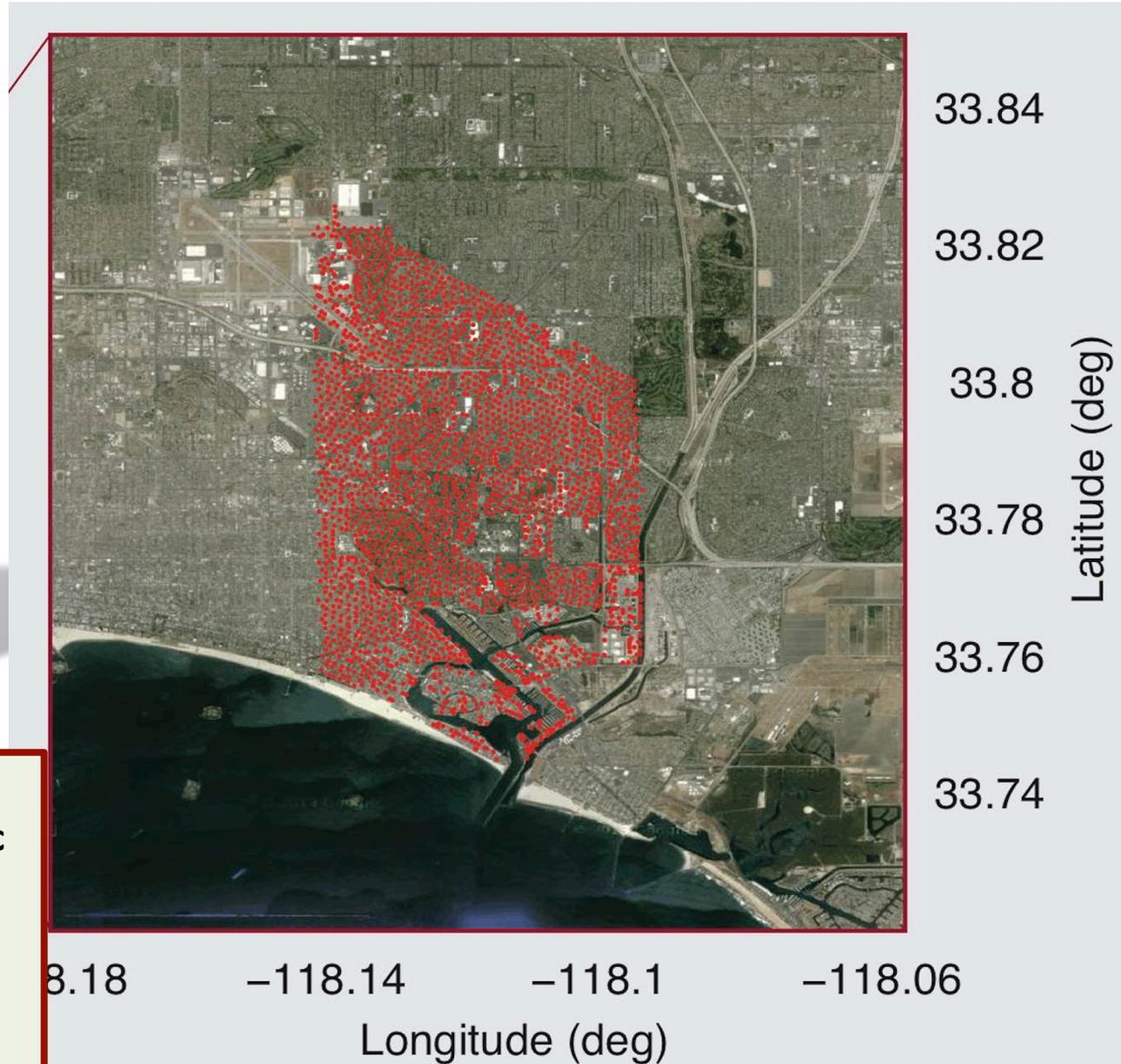
Courtesy of Pierre Boue'

Value of Dense Recording



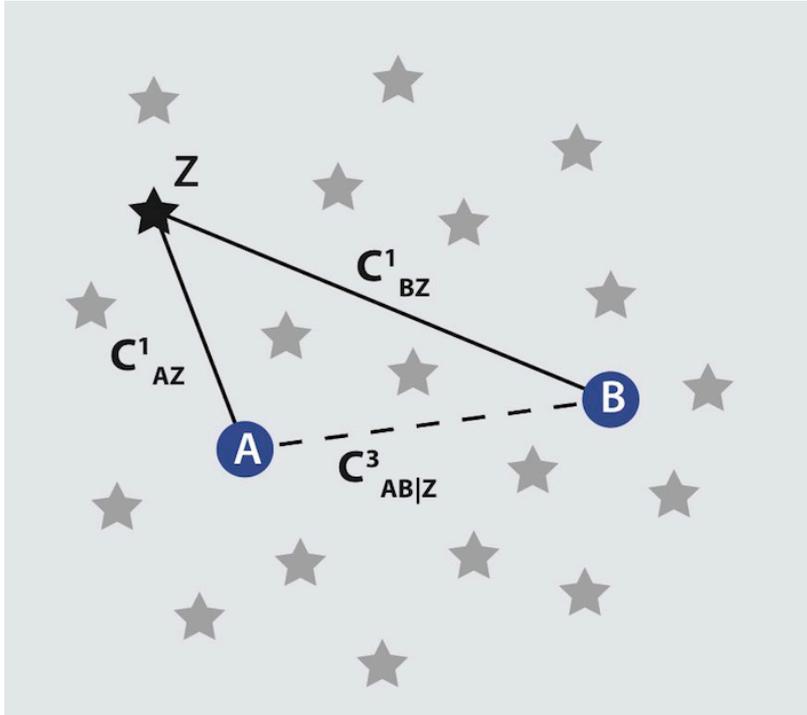
Data provided by Signal Hill
Petroleum and NodalSeismic

~2500 vertical Sensors
~100 m Spacing
#obs = $N(N-1)/2$



Correlation of Correlations

(Stehly et al., 2008)

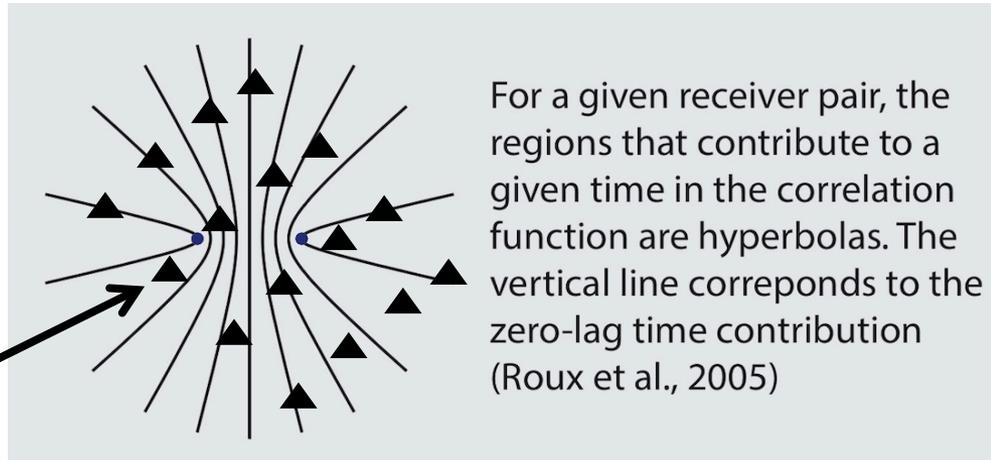


Correlation of Green's functions derived from correlations

Approximates equi-partitioned ideal

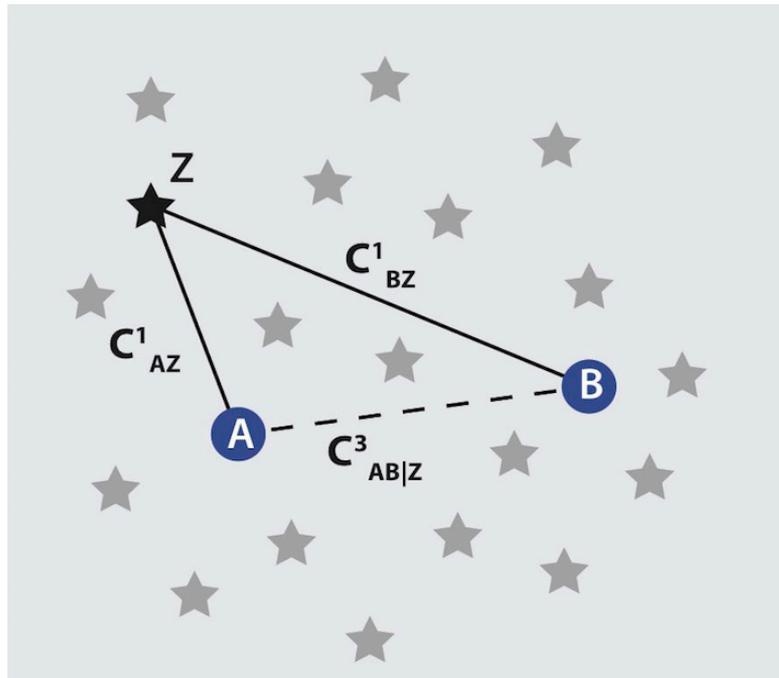
Can compare theoretical and observed kernels to understand and remove amplitude bias

Map this contribution using the virtual sources at each station in Long Beach array.



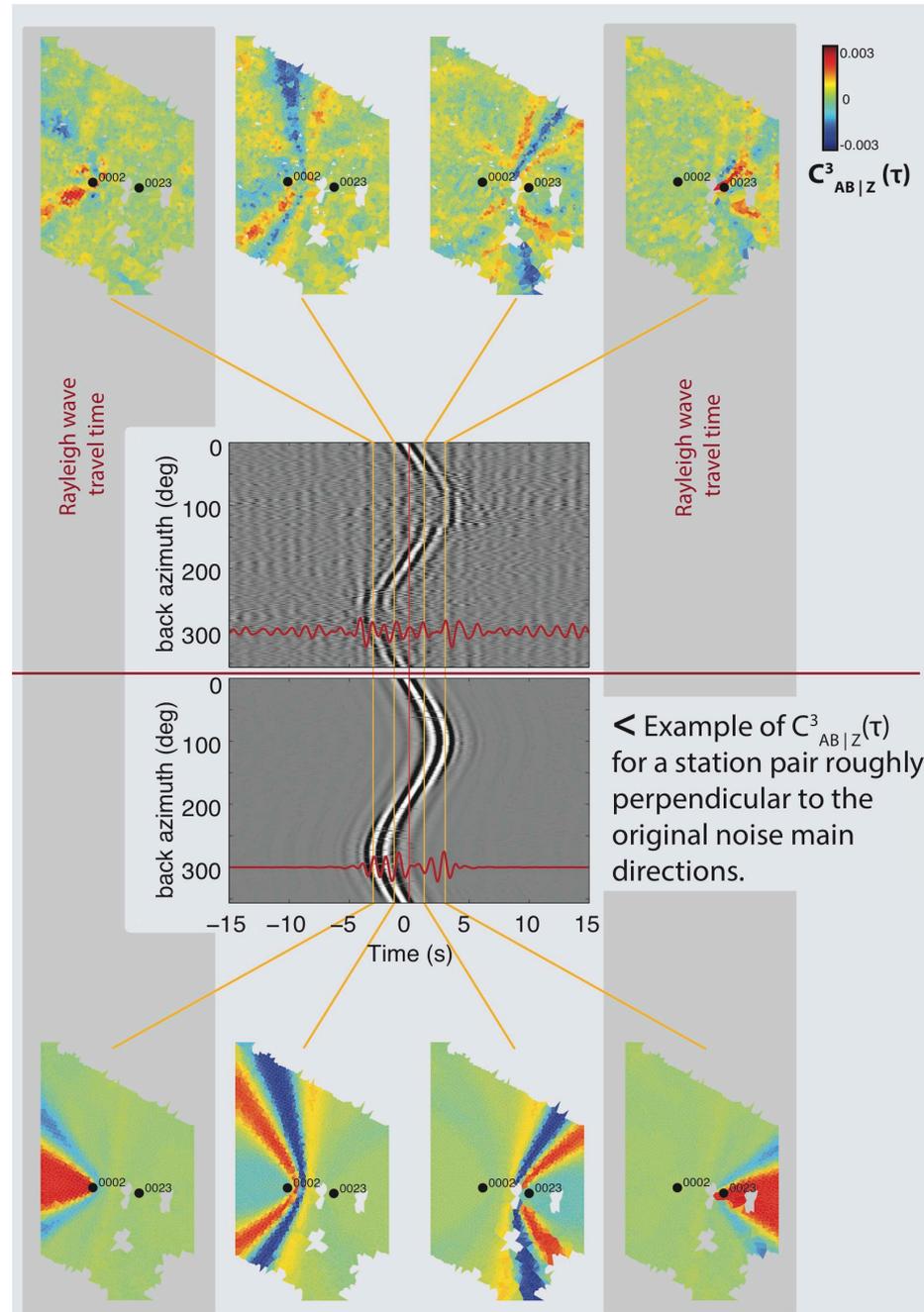
For a given receiver pair, the regions that contribute to a given time in the correlation function are hyperbolas. The vertical line corresponds to the zero-lag time contribution (Roux et al., 2005)

Correlation of Correlations

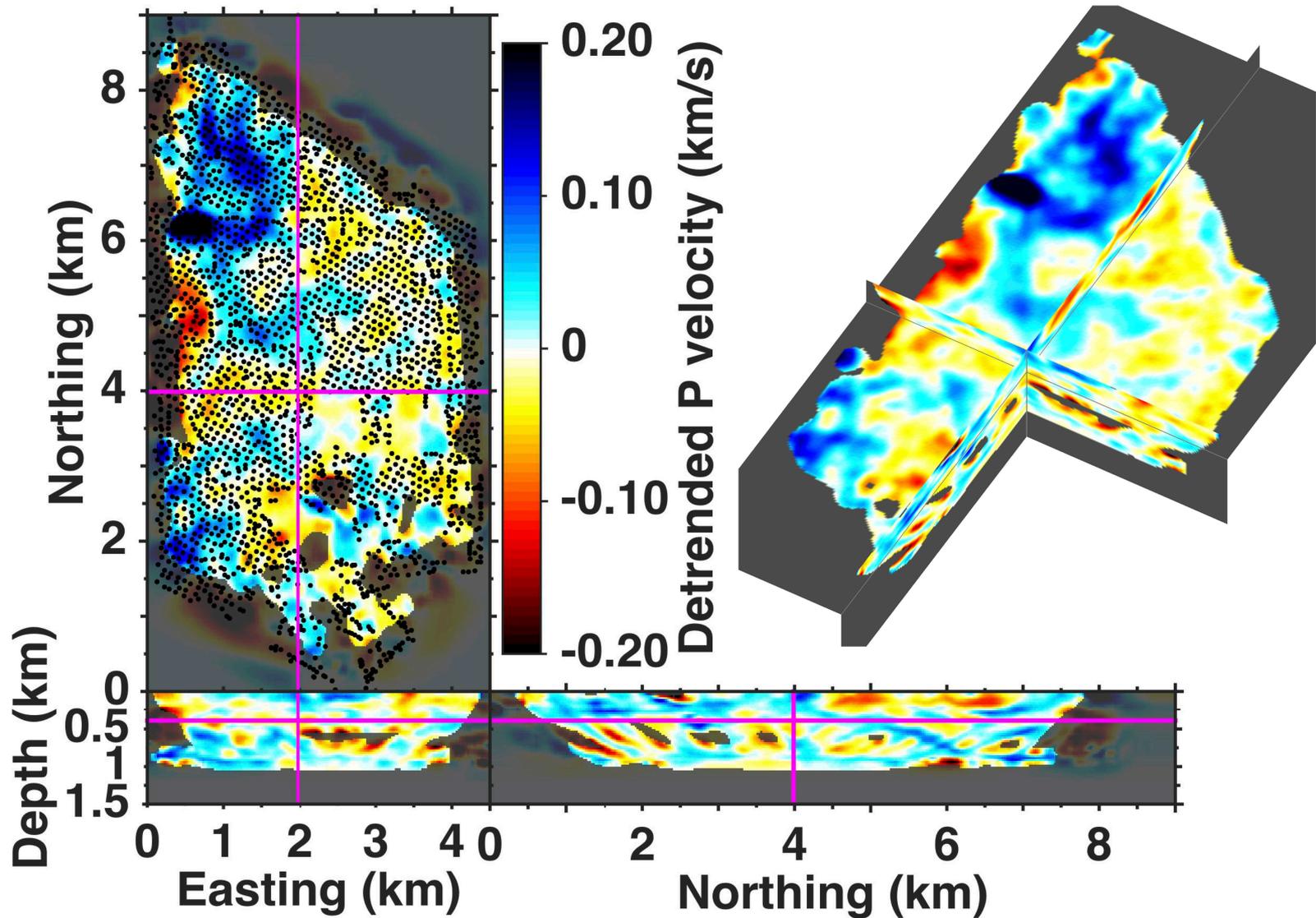


Can correct for remaining bias from lack of equi-partitioning

With 3 component data could correct for excitation bias as well



3D Body-Wave Tomography with ASF



Conclusions – ASF for GMP

Caveats:

1. Only part of the problem (source, nonlinearity)
2. Bias (azimuthal, multi-component)
3. Bandwidth limitations
4. Body waves

Promise:

1. New approaches to an important problem
2. Active, not passive
 - no need to wait for real earthquakes
 - can sample sources/sites of strategic interest
 - can revisit past earthquakes
3. Resolves 3D effects – Earth does the experiment

An Assertion

GMP prediction for IS is a solvable problem

- we know where to put the instruments
- small, frequent earthquakes are important
- we can trigger the earthquakes
- reducing epistemic uncertainty and characterizing variability is *much* easier than for tectonic events

A Recommendation

Dense 3-component instrumentation of IS sequence(s)

- Don't need to record very long for high frequencies
- Can gather critical data even without earthquakes
- If earthquakes occur then a wealth of data for GMP